

ERTH2404

Lecture 3: Sedimentary Rocks



Photos: M.Lamontagne

Red Rock Canyon State Park, CA

Maurice Lamontagne

Reading assignment

- Please read Kehew's book to complement the material presented in this lecture:
Chap. 5;

Lecture content

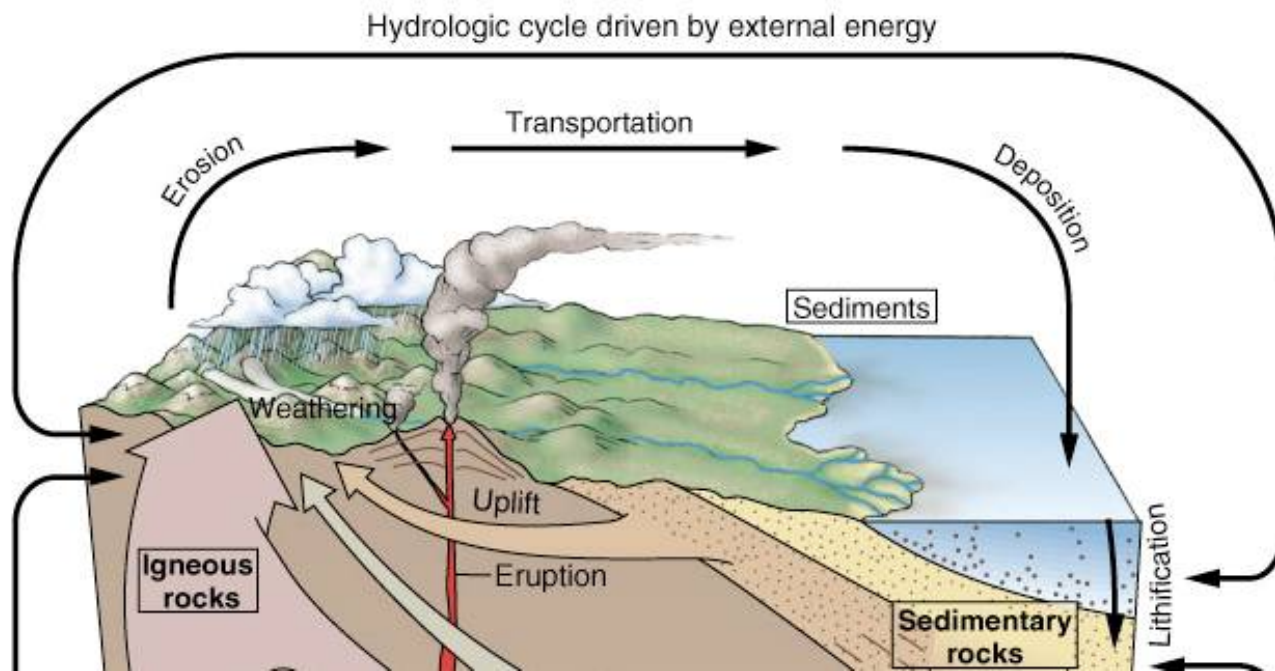
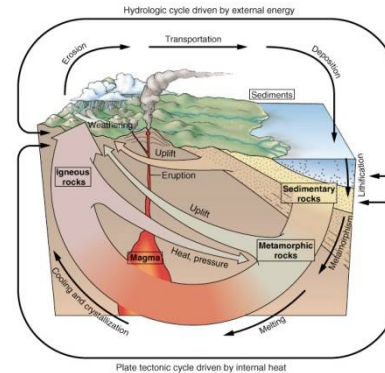
- Origin and classification of sedimentary rocks
 - Detrital sedimentary rocks
 - Chemical sedimentary rocks
- Sedimentary structures
- Sedimentary depositional environments
- Engineering considerations

Sedimentary Rocks

- Significance
 - Sedimentary rocks make up 5% of crustal rocks
 - 75% of rock exposed
 - Important economically
 - Building material
 - Energy sources (coal, petroleum, natural gas)
 - Sources of iron, aluminum and manganese
 - Record Earth's past history and environments

Remember the rock cycle

What happens to rocks at the Earth's surface?



Remember the rock cycle

- Sedimentary rocks are formed from any type of pre-existing rocks by the processes of
 - Weathering
 - Erosion & Transportation
 - Deposition
- Common characteristic: BEDDING (layering)
- Sedimentary rocks reflects the environment of erosion (climate, ocean vs. land)

Classification of Sedimentary Rocks

- Two criteria:
 1. Process of formation: detrital or chemical
 2. Texture: grain size, shape, physical features
- **Detrital (clastic)**: formed by the erosion and deposition of rock fragments
- **Chemical**: formed by precipitation

Texture

- **Clastic rocks** are composed of aggregates of individual fragments:
 - **Clasts** (larger fragments) surrounded by a **matrix** (fine-grained sediments)
 - Commonly layered
- **Non-clastic rocks** are composed of a pattern of interlocking crystals
 - Similar to crystalline igneous rocks

Classification Table

Process of formation		Texture	
Detrital		Clastic	
Chemical	Inorganic		Non-clastic
	Biochemical		

Process of Formation

- Detrital Sedimentary Rocks
 - **Weathering**: breakdown of rocks into fragments
 - Physical (abrasion, freezing), chemical
 - **Erosion**: Detachment and transport of fragments by water, wind or gravity
 - **Deposition**: sediments deposited by natural process
 - **Lithification**: cementation and compaction
- Chemical Sedimentary Rocks
 - **Precipitation**: formation of a solid in a solution during a chemical reaction

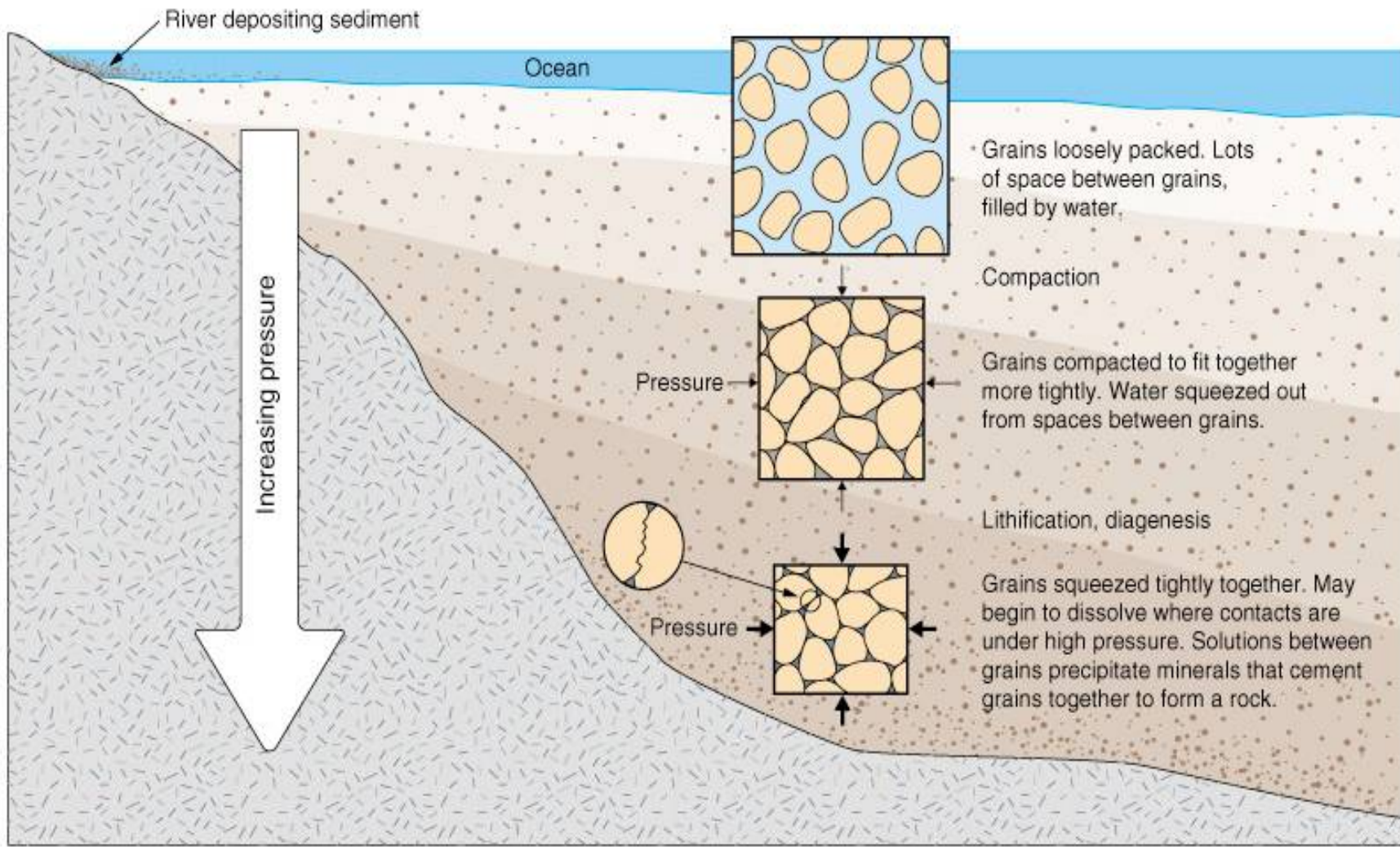
Lithification

- **Lithification** (process in which sediments compact under pressure, expel connate fluids, and gradually become solid rock)
- Achieved in several ways
 1. Compaction: caused by weight of overlying material
 - Most basic lithification process
 - Mechanical reduction of void space in between particles
 2. Cementation: filling of void spaces by chemical precipitation
 - Generates cement that binds particles together
 - Two most common cements: quartz (SiO_2) and calcite (CaCO_3)

Lithification

- Achieved in several ways
 - 3. Crystallization: additional growth of the original crystal within the void spaces
 - Without addition of new chemical substances
- Impact of lithification
 - Reduces the amount of void space between particles
 - Increase in density, decrease in porosity & permeability
- The degree of lithification influences the induration of sedimentary rocks

Lithification



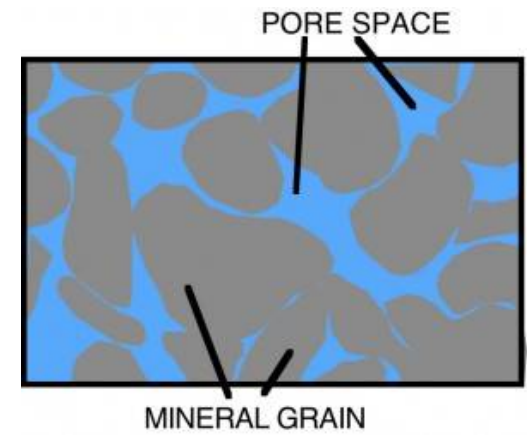
Porosity

- Total volume of Earth material: V_T
Total volume of void space: V_v
Total volume of particles: V_s

- **Porosity** [%] $\eta = V_v / V_T$

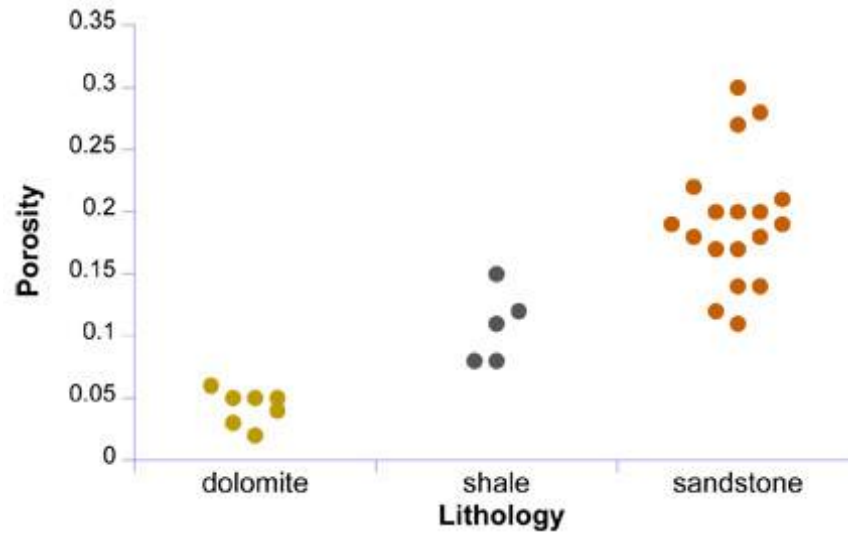
- **Void ratio** [%] $e = V_v / V_s$

- Note: porosity is not the same as permeability...

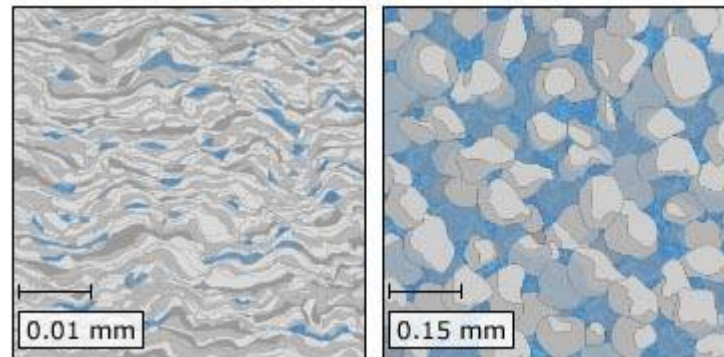


Source: Wisconsin
Geological and
Natural Survey

Some porosity values



Source: Wisconsin Geological and Natural Survey



Shale

Sandstone

Permeability

- **Permeability**: general qualitative term describing the ability of a material to transmit fluids
- Permeability is a composite property of:
 - Material properties: size, shape and interconnectivity of the voids
 - Fluid properties: density, viscosity
- Permeability varies by several orders of magnitude in earth materials

Classification of detrital rocks

- Most common minerals are quartz, feldspar, clay minerals
- Detrital rocks are classified according to **particle/grain** size

Classification of detrital rocks

- Wentworth scale

Grain diameter [mm]	Size name
<1/256	clay
1/256 - 1/16	silt
1/16 - 2	sand
>2	gravel

Detrital rocks in the field

- http://www.glossary.oilfield.slb.com/Terms/u/udden-wentworth_scale.aspx
- Large blocks are found near the source whereas very fine grains end up in deep marine environments far from the source



Lake Powell, UT Photo: M.Lamontagne

Classification of detrital rocks:

Grain diameter [mm]	Size name	Detrital sedimentary rock
<1/256	clay	Mudstone
		Shale
1/256 - 1/16	silt	Siltstone
1/16 - 2	sand	Sandstone Arenite Wacke
>2	gravel	Conglomerate
		Breccia

Properties of detrital rocks

- **Grain size:** based on grain diameter
- **Roundness:** rounded vs. angular
- **Sorting:** distribution of grain sizes in sedimentary rock

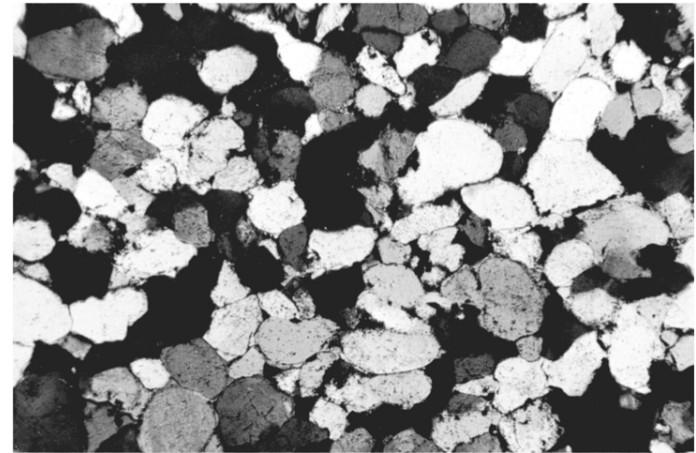
Grain size, Roundness, Sorting

- Grain size is related to the distance traveled from the **sediment** source
 - Due to abrasion
 - Mineral composition: how tough is a mineral?

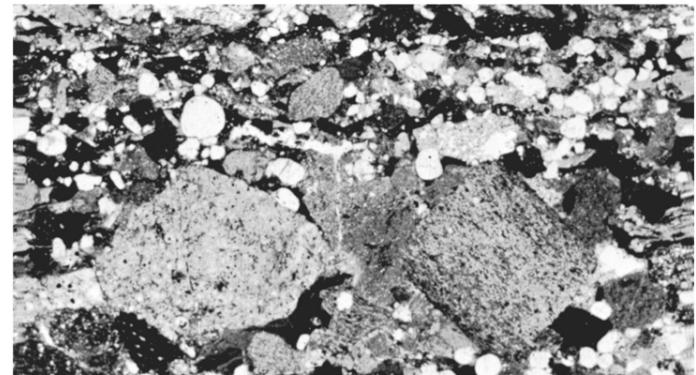
	Close to Source	Far from Source
Grain Size	Larger	Smaller
Roundness	Angular	Rounded
Sorting	Poorly sorted	Better sorted

Sediment Maturity

- **Mature** vs. **Immature**
- Mature: well sorted, fine-grained (sand or smaller), rounded, mostly quartz grains
- Immature: poorly sorted, coarse-grained, angular, very little proportion of quartz



(a)



(b)

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Common detrital rocks

- Uniform grain size
 - **Shale**: composed of mud
 - **Siltstone**: composed of silt
 - **Sandstone**: composed of sand (gritty)
 - Arenite (< 15% matrix) (from latin: arena: sand)
 - Wacke (> 15% matrix)

Common detrital rocks: Shale

- Most common sedimentary rock (50% of all sed. rocks)
- Formed by the compaction and/or cementation of clay minerals
- Particles ($< 1/256$ mm) deposited in thin layers (laminellae)
- Shale splits into thin layers (fissility)
- Usually parallel to depositional layering
- Shale weathers readily
- Found in gentle slopes
- Important source of hydrocarbons (shale gas, shale oil)

Common detrital rocks: Sandstone

- Sandstone is formed in a variety of environments
- Texture can be used to interpret the rock's history
- Quartz (SiO_2) is the prominent mineral
- More durable than shale
- Found in steeper slopes

Common detrital rocks: Sandstone

- Quartz Sandstone:
 - All sand-sized, all rounded, well sorted, 99% quartz
 - monomineralic rock





Lake Powell, UT Photo: M.Lamontagne



Bryce Canyon, UT Photo: M.Lamontagne



- Arches NP, UT Photo: M.Lamontagne

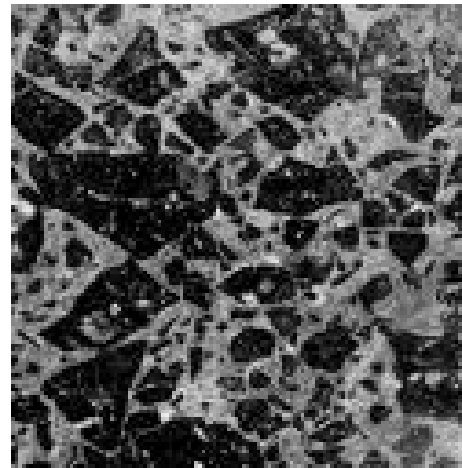


- Arches NP, UT Photo: M.Lamontagne

Common detrital rocks

- Wide range of grain size > 2 mm
 - **Breccia**: composed of angular gravel
 - Deposited closer to source
 - **Conglomerate**: composed of rounded gravel
 - Transported further from source

conglomerate



breccia

USGS

Chemical sedimentary rocks

- We now discuss chemical sedimentary rocks



Great Salt Lake, UT Photo: M.Lamontagne

Chemical sedimentary rocks

- Material precipitated from water
 - Inorganic: minerals crystallize from supersaturated water
 - Biochemical/Organic: organisms extract compounds from seawater to form shell; when dies, sinks to bottom of ocean
 - Biochemical sediment

Inorganic Precipitation

- Water supersaturated in certain elements
 - Often because of evaporation
- Precipitated crystals sink to bottom
- Ca, CO₃ → limestone (CaCO₃)
- Ca, Mg, CO₃ → dolostone (CaMg(CO₃)₂)
- SiO₂ → chert
- Na, K, Cl → rock salt (NaCl, KCl)
- Ca, SO₄ → gypsum (CaSO₄ · 2H₂O)

Inorganic Precipitation

- SiO_2 based
 - Chert, opal ($+\text{H}_2\text{O}$): microcrystalline
 - Geode: mineral deposits in cavities in rocks



Chert



Geode

Biochemical/Organic Precipitation

- Organisms extract minerals from seawater and build shells
- Shells sink when organisms die
- Reef, benthic and some planktonic organisms extract Ca and CO_2 → limestone (CaCO_3)
- Other planktonic organisms extract SiO_2 → chert

Limestone

- Most common chemical sedimentary rock (10% of all sedimentary rocks)
 - Composed mainly of mineral calcite (CaCO_3)
 - Produced by **biochemical** and **inorganic** precipitation
- Marine biochemical origin most common
 - Coral reefs (important petroleum reservoir rock)
 - Coquina (broken shells)
 - Chalk (shells of microscopic organisms)

Generalized bedrock geology,
Ottawa-Hull, Ontario and Quebec; Harrison, J E; MacDonald, G.
Geological Survey of Canada, "A" Series Map 1508A, 1979, 1 sheet,
<https://doi.org/10.4095/109223>

- The map can be downloaded at:
- <https://geoscan.nrcan.gc.ca/starweb/geoscan/servlet.starweb?path=geoscan/download.web&search1=R=109223>

Limestones in the Ottawa-Gatineau Region



Photo: GeoheritageOttawa



- Photo: NRCan

Organic limestone

- Coquina: composed of poorly cemented shells



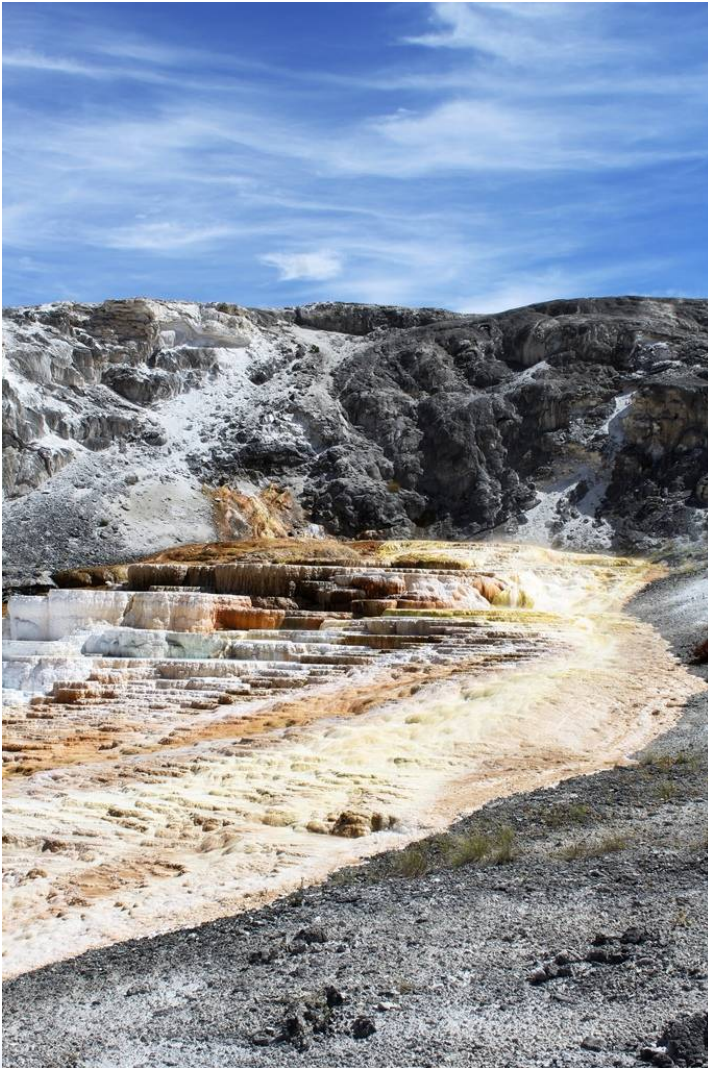
Scale in mm

Inorganic limestone

- Travertine (limestone): stalactites, stalagmites, hot springs



Inorganic limestone



- Travertine deposits: Hot springs (Yellowstone)

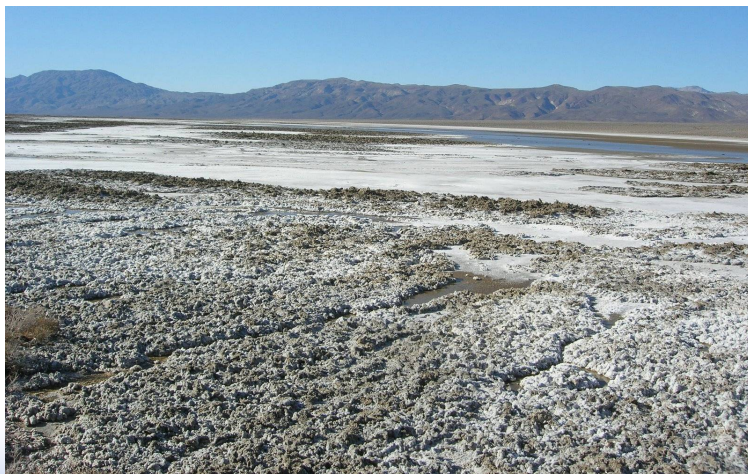
Photo: M. Lamontagne

Inorganic Precipitation

- Evaporites
 - Precipitation due to evaporation of shallow lakes, inland seas
 - Minerals precipitate in order of increasing solubility:
 1. $\text{Ca} + \text{SO}_4 \rightarrow \text{gypsum (CaSO}_4 \cdot \text{H}_2\text{O)}$
 2. $\text{Na} + \text{Cl} \rightarrow \text{halite (NaCl)}$
 3. $\text{K} + \text{Cl} \rightarrow \text{sylvite (KCl)}$

Salt flats

- Playa (ephemeral) lake: a waterbody existing for a short period following precipitation or snowmelt
- In a desert environment where rain dissolves elements from rocks, washes down into valleys ephemeral lakes
- Salt flat: water evaporates leaving behind salts



Credit: Eos, AGU

Salt flats

- Bonneville Salt Flats (Utah)



Sedimentary structures

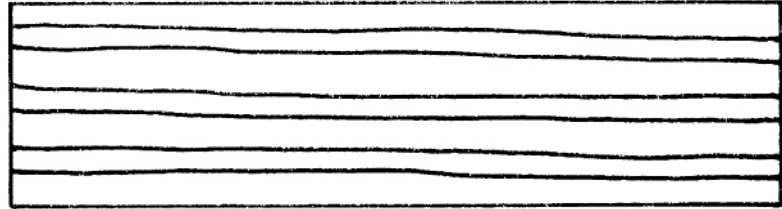
- Sediments deposited in layers: beds, **strata**
- Each layer unique, records conditions under which particles, minerals deposited
- **Bedding plane**: surface that separates each successive layer
- Conditions recorded in **sedimentary structures** within layers
- **Folding**: strata deformed and curved

Sedimentary structures: Bedding

- **Parallel bedding:** quiet deposition of particles falling to the bottom of a water column by gravity
- **Cross bedding:** deposition under alternating wind or water current direction
- **Graded bedding:**
 - Coarse particles at the bottom, changing to finer particles at the top
 - Deposition of material moving rapidly downslope

Sedimentary structures: Bedding

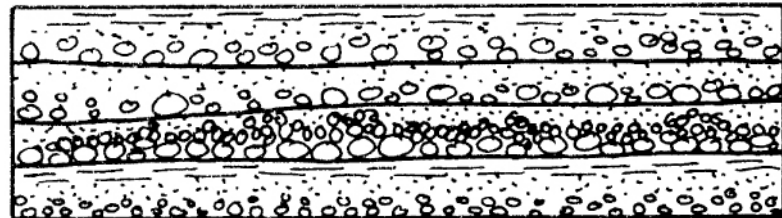
- **Parallel bedding**



- **Cross bedding**

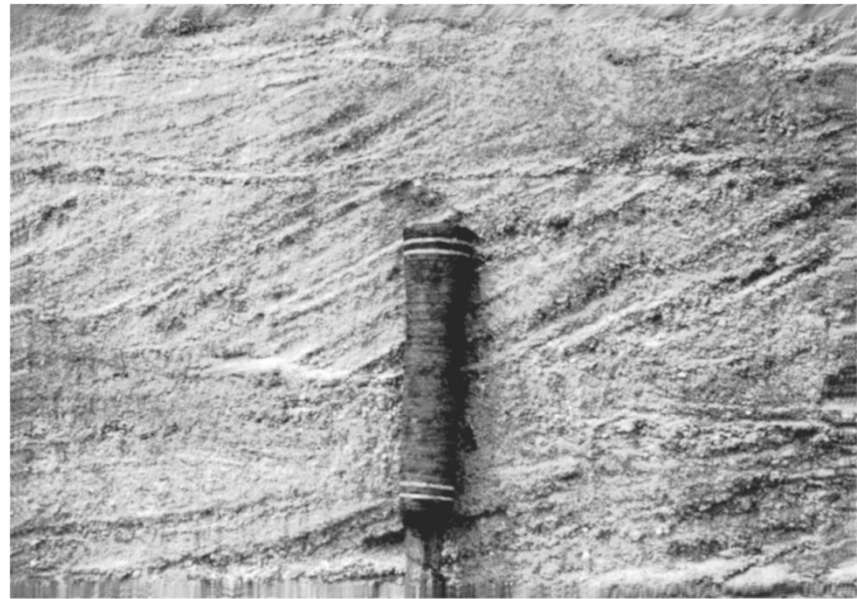


- **Graded bedding**



Sedimentary structures: Cross-bedding

- Sand dunes, some river deposits
- Records alternating wind or current direction



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Sedimentary structures: Ripples

- **Ripple marks**
 - Small scale back-and-forth motion of waves
 - Occur on beaches and streams



Sedimentary structures: Mud cracks

- Muddy sediment dries and shrinks
 - cracks fill with different sediment (desert, dry lakes)



Fossils



- Remains of plants or organisms
- shells, bones, petrified wood, imprints of plants
- **index fossils**: indicator of relative time

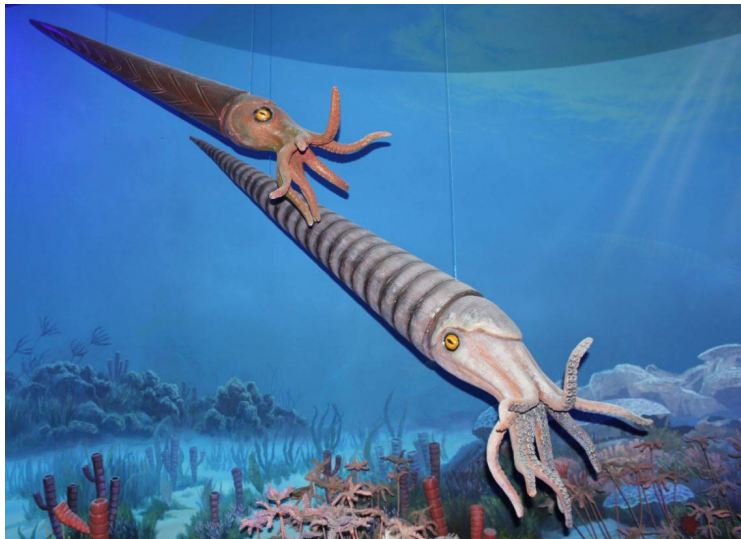


Photo: M. Lamontagne
Museum of Natural Sciences, Beijing

Sedimentary environments

- Features of sedimentary rocks reflect their depositional environments
- Important factors:
 - Transporting agent (water, wind, ice)
 - Flow characteristics
 - Characteristics of body of water
 - Size, shape, depth, circulation of water
 - Presence of life
 - Nature of sediments being deposited

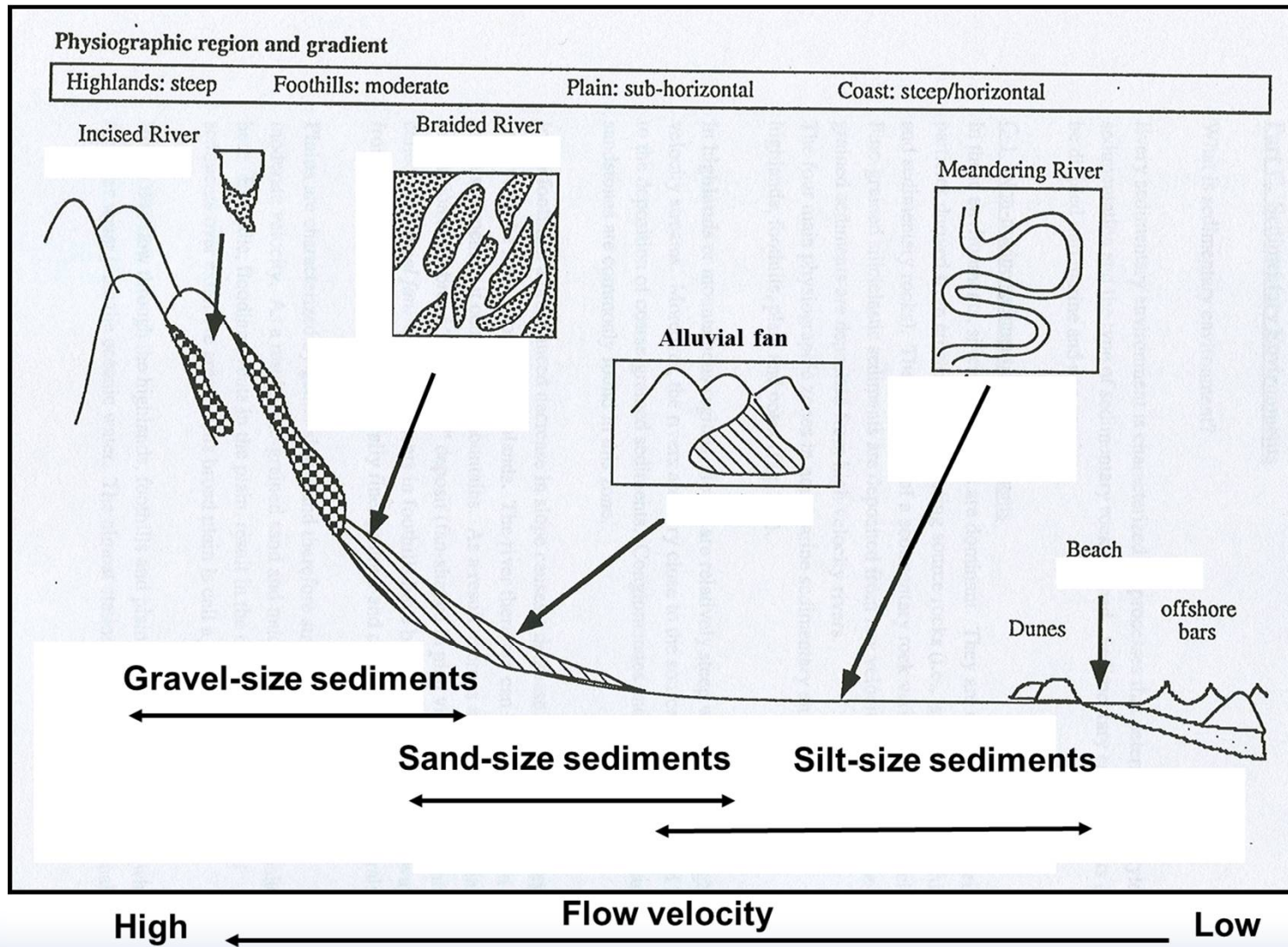
Sedimentary environments

- Types of sedimentary environments
 1. Continental
 2. Marine
 - Shallow (continental shelf; depth \leq 200 m)
 - Deep (seaward of continental shelf)

Continental environments

- Dominated by erosion and deposition associated with water, ice and wind
- Detrital sedimentary rocks dominant
- Grain size varies with flow velocity
 - Coarse-grained sediments deposited by fast-flowing rivers
 - Fine-grained sediments deposited by slow-flowing rivers

Continental environments



Great Sand Dune NP, CO



Photo: M. Lamontagne

Marine environments

- Shallow
 - Detrital sedimentary rocks
 - Near shore:
 - Low-energy environment (e.g. protected lagoon): silt-size sediments
 - High-energy environment (e.g. exposed beach): sand-size sediments
 -
 - Shelf: wide range of sediment sizes
 - Slope: frequent mass movements, graded bedding

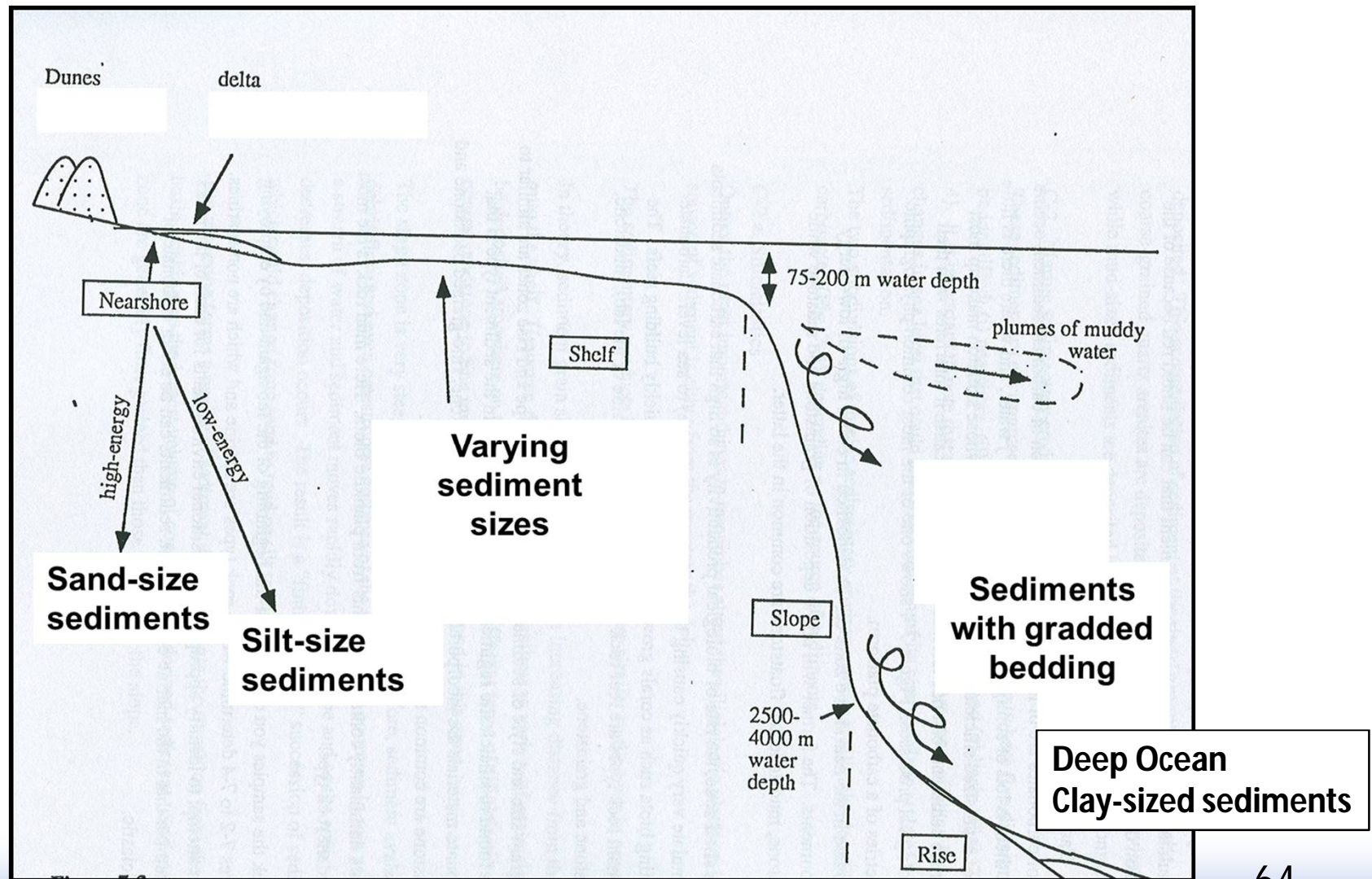
Marine environments

- Shallow
 - Biochemical sedimentary rocks
 - Conditions for accumulation of biochemical sedimentary rocks
 - Sunlight
 - Oxygen
 - Warm water ($T > 20^{\circ}\text{C}$)
 - Little silt- and clay-size sediments

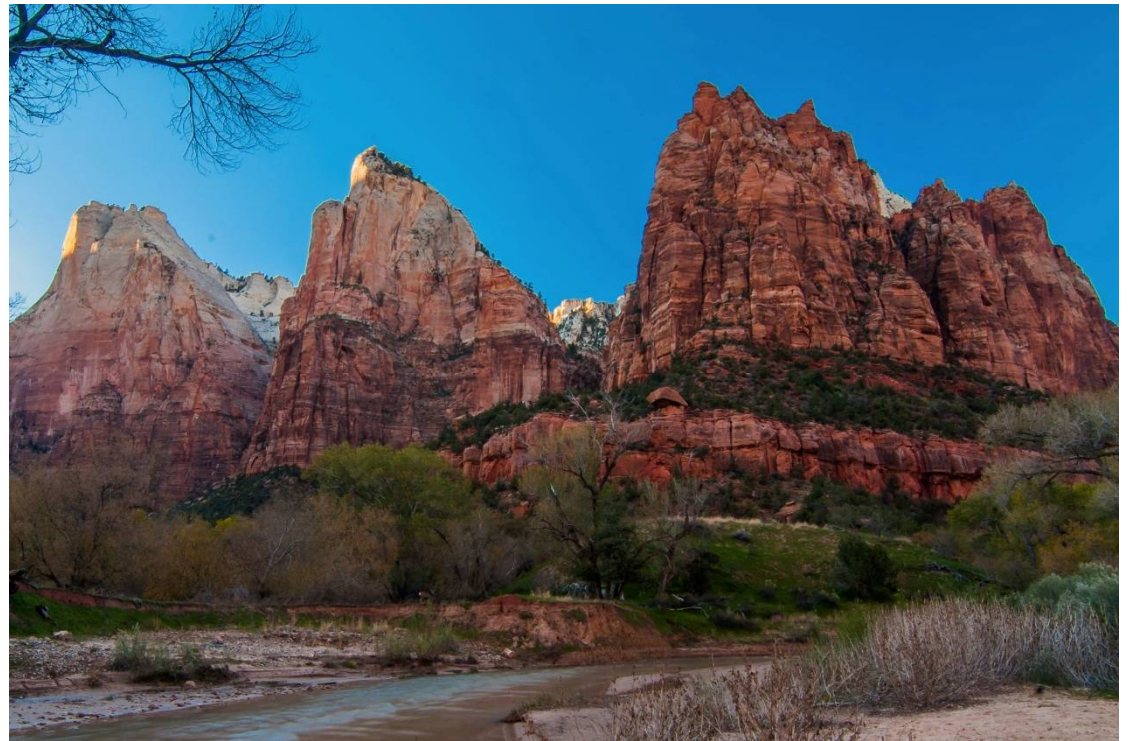
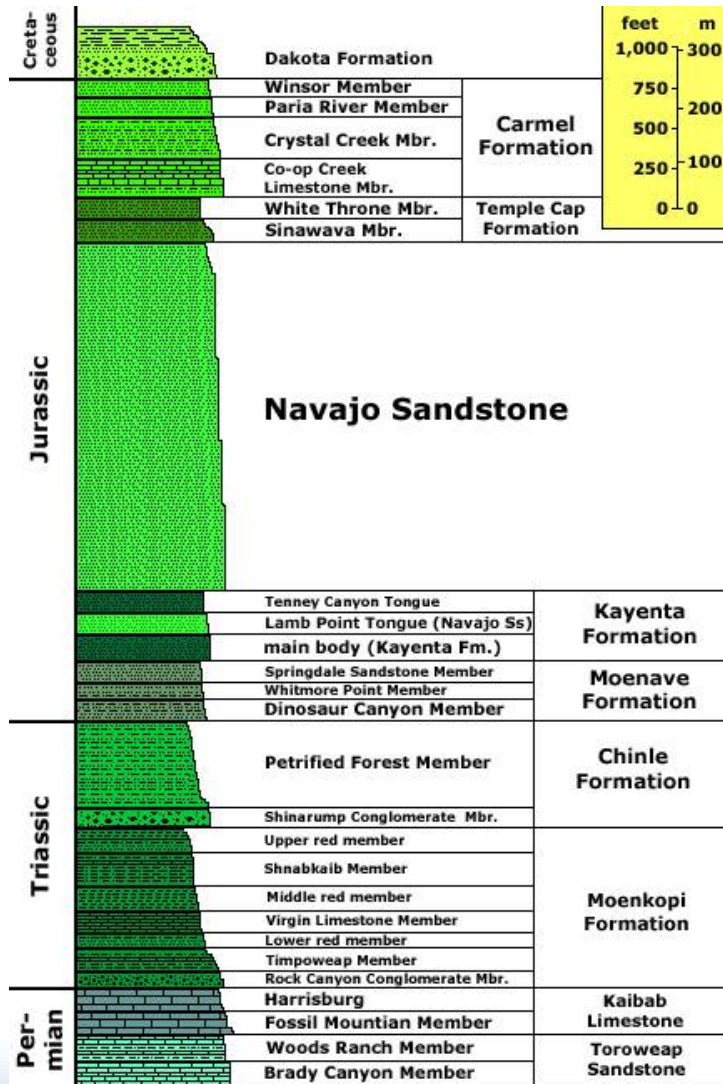
Marine environments

- Deep ocean
 - Clay-size sediments
 - Quiet environment
 - Conducive to the formation of shale
 - No biochemical sedimentary rocks
 - Reefs cannot develop

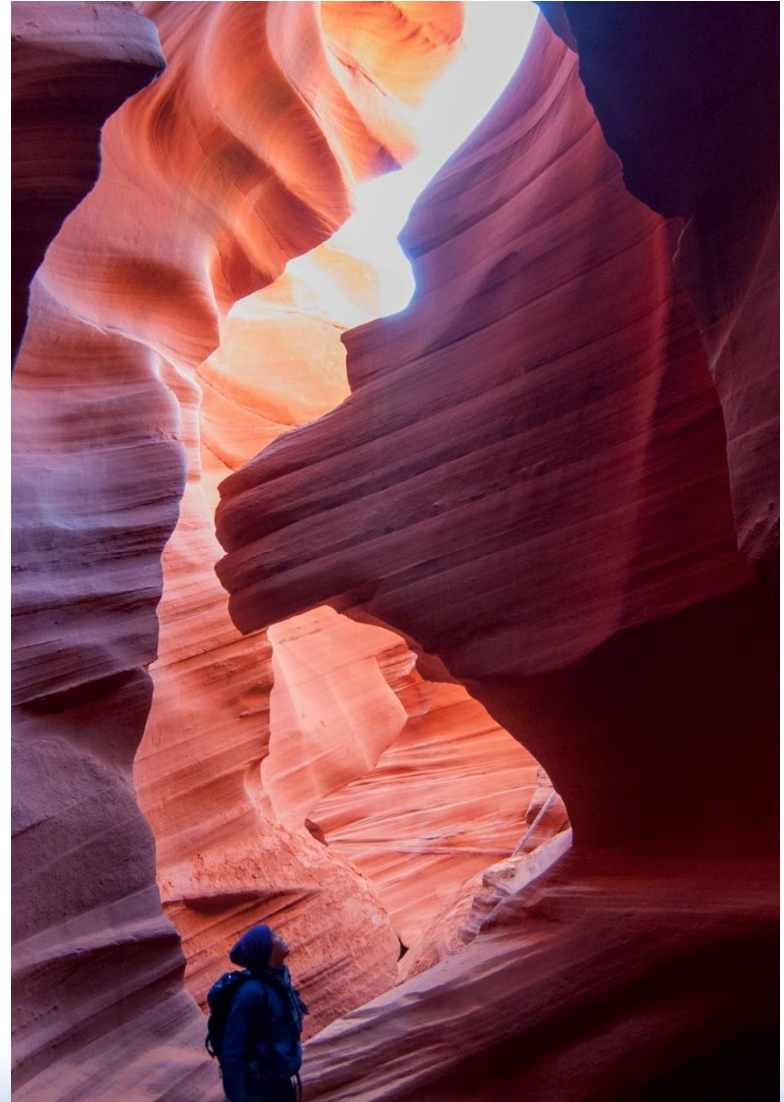
Marine environments



Zion National Park: Navajo Sandstone



Navajo Sandstone: Antelope Canyon



Dinosaur Provincial Park, Alberta

- World Heritage Site since 1979
- Mesozoic era (Cretaceous (≈ 75 Ma))
- Warm inland sea at the base of Cordillera
 - Subtropical climate
 - Lush vegetation in coastal plain
 - Rivers carried organic-rich sediments into sea
- Layer-cake sedimentary deposits
 - Mostly shale, sandstone and mudstone
 - Rich deposit of dinosaur fossils

Dinosaur Provincial Park, Alberta

- Valleys carved by water from melting glaciers at the end of the last glaciation ($\approx 13,000$ years ago)



Photo: C. Samson, Carleton U.

Dinosaur Provincial Park, Alberta

- Formation

where the prairie landscape suddenly drops away, exposing millions of years of history before you.

68 million years ago... the Rocky Mountains were beginning to rise and the shallow Bears paw Sea lay to the east. Large rivers wound their way across the land, threading through ancient forests of cypress and dawn redwood. Along their journey, these rivers deposited sediment. Sands were left in the channels and fine silts and clays were dropped along the floodplains and in calm backwaters. Today, these layers of light-coloured sandstone and darker siltstone lie before you. In some areas of the valley they are separated by seams of coal remnants of once vast swamps. Throughout the valley you can even spot the remains of some of the many coal mines which supported the economy of Drumheller and nearby communities.

Photo: C. Samson, Carleton U.

Engineering considerations

- Water infiltrates porous/permeable rocks
- Cracks open after several freeze/thaw cycles

Engineering considerations: Detrital

- Variability
 - Vertical and horizontal gradation into other sedimentary rocks
 - Variations within each strata
 - Knowledge of depositional environments may allow for prediction of rock type and properties
 - Quarries, exploration areas, oil and gas exploration

Engineering considerations: Detrital

- Bedding causes extremes in directional properties
 - **Anisotropy**: characteristic of a property having a different value when measured in different directions
 - Strength, permeability, seismic velocity
- Bedding planes are zones of weakness
 - Exploit them in quarrying

Engineering considerations: Detrital

- Engineering properties influenced by geological events occurring long after sediment deposition
 - Lithification affects strength
 - Strength influenced by type and amount of cement
 - Quartz (SiO_2) cement stronger than calcite (CaCO_3) cement
 - Conglomerates often weak due to poor cementing between clasts and matrix

Engineering considerations: Chemical

- Carbonate rocks: limestone and dolostone
- PROS
 - Best sedimentary aggregates
 - Excellent source of building rock
 - Carved limestone blocks used as foundation in older Ottawa neighborhoods
- CONS
 - Thin bedded limestones separated by layers of clay or shale may serve as sliding planes
 - May dissolve and form a system of subterranean galleries (karst terrain)

Next: Metamorphic Rocks



ERTH2404

Metamorphic Rocks



Gneisses

Maurice Lamontagne

Reading assignment

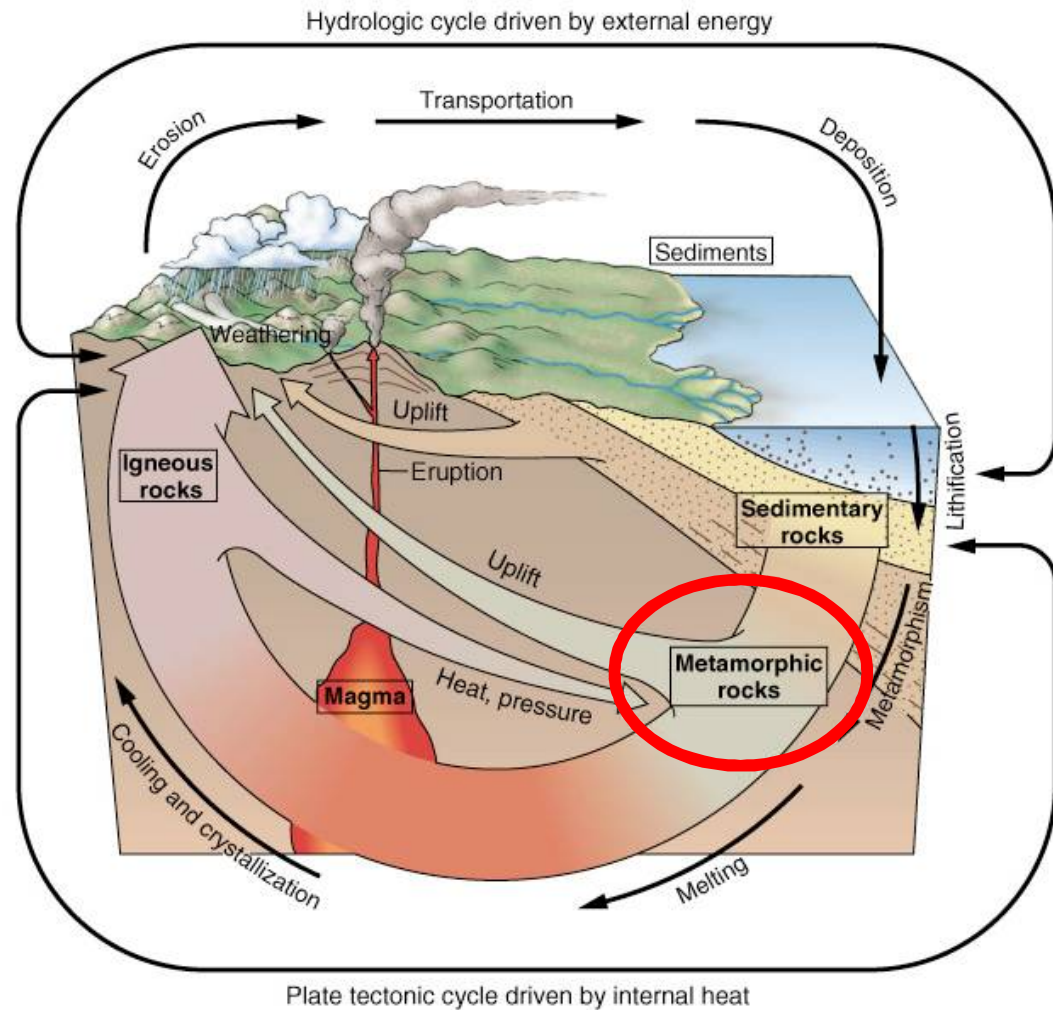
- Please read Kehew's book to complement the material presented in this lecture:
Chap. 6;

Lecture objectives

- To learn the different types of metamorphism
- To understand the relationship between metamorphic rocks and other rock types
- To be able to characterize and identify metamorphic rocks
- To be aware of how the characteristics of metamorphic rocks might impact engineering work

Remember the rock cycle

- Metamorphic rocks are formed from any type of pre-existing rock



Metamorphism

- **Metamorphism**: transformation of one rock into another through the action of
 - Heat
 - Pressure
 - Chemically active fluids (metamorphic agents)
- Rocks retain their solid form
 - But they can flow in a plastic-manner (i.e very slow deformation)

Metamorphism

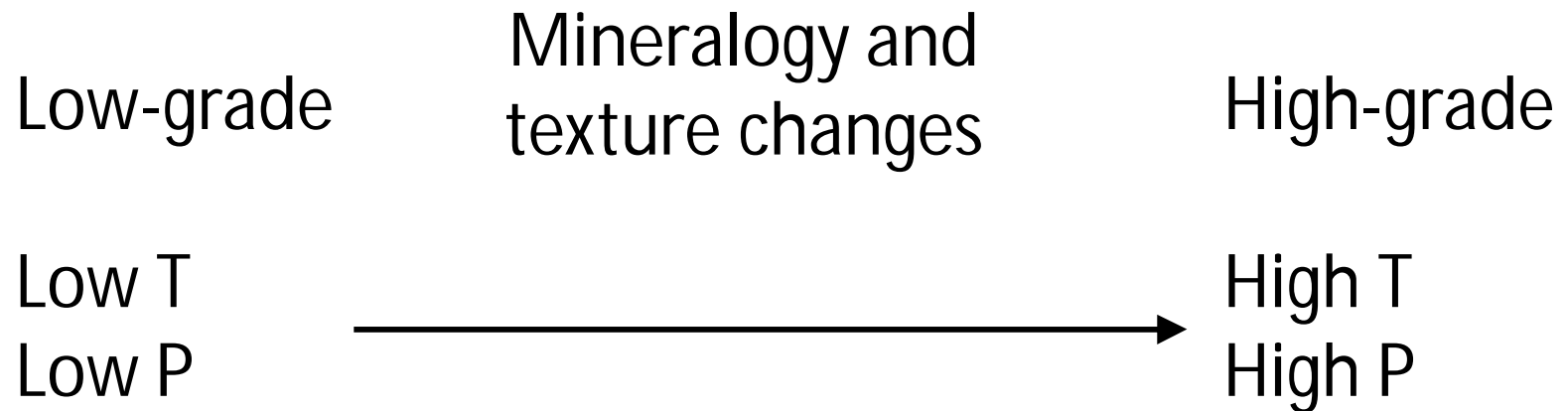
- Metamorphism may cause many changes to rock
 - Formation of larger or smaller grains
 - Reorientation of mineral grains
 - Texture such as banded appearance, sheen
 - Increase in density
 - Transformation of existing minerals into minerals stable in the new T-P conditions

Metamorphism

- Metamorphism occurs in the Earth's crust over 100s of Ma
 - In the core of mountain belts
 - At depth within stable continental interiors (shields)
 - Often in association with igneous intrusions
- Metamorphic rocks can become exposed at the surface when overlying rocks have been removed by erosion or when tectonic movements bring them closer to the surface
 - Exposed outcrops are commonly of Precambrian age due to the long time needed for erosion to occur

Metamorphism

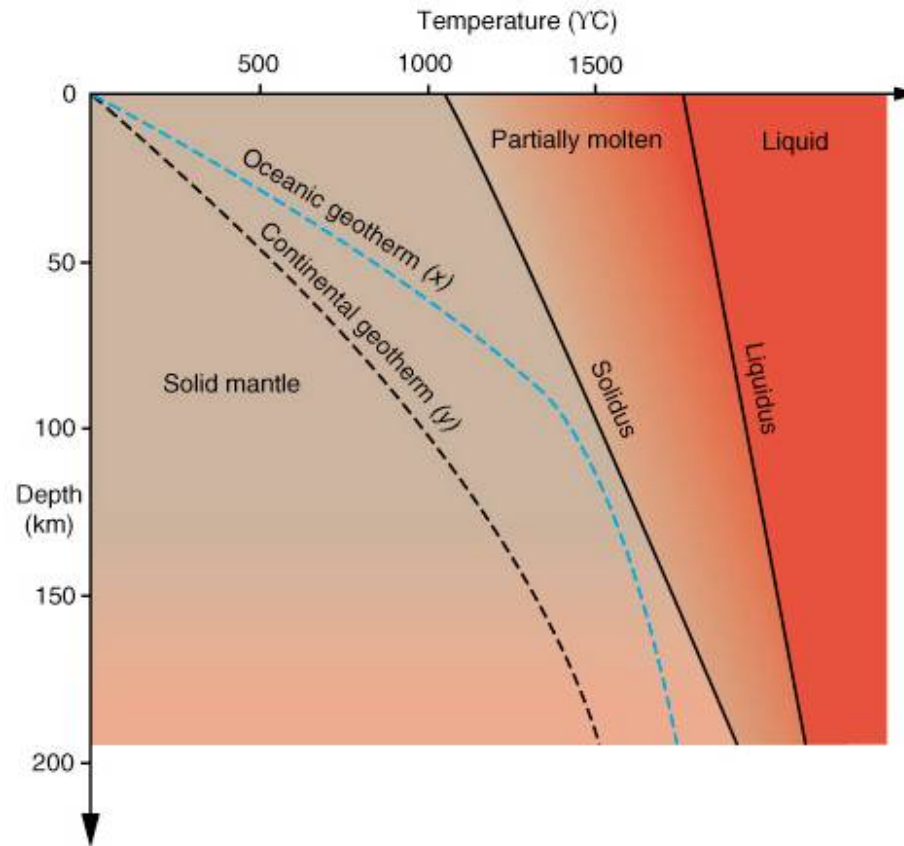
- **Metamorphic grade**
 - Certain minerals and textures are diagnostic of the degree of metamorphism experienced by rocks



Metamorphism: Temperature

- Most important factor for metamorphism
- Drives chemical changes
 - Recrystallization of existing minerals into new minerals
- Impacts:
 - Texture, crystal structure, crystal size
- Two sources:
 - Magma intrusion
 - Geothermal gradient (increase of T with depth)

Metamorphism: Temperature



Geothermal gradient

Metamorphism: Pressure

- Approximate increase of 1 kbar per depth of 4.4 km
- **Confining pressure:** pressure applied in all directions
 - Causes decrease in rock size and an increase in density
 - Does not fold or deform rock
- **Differential stress:** pressure applied from specific directions
 - In ductile regime, can deform rocks
 - Associated with mountain building and faulting

Metamorphism: Fluids

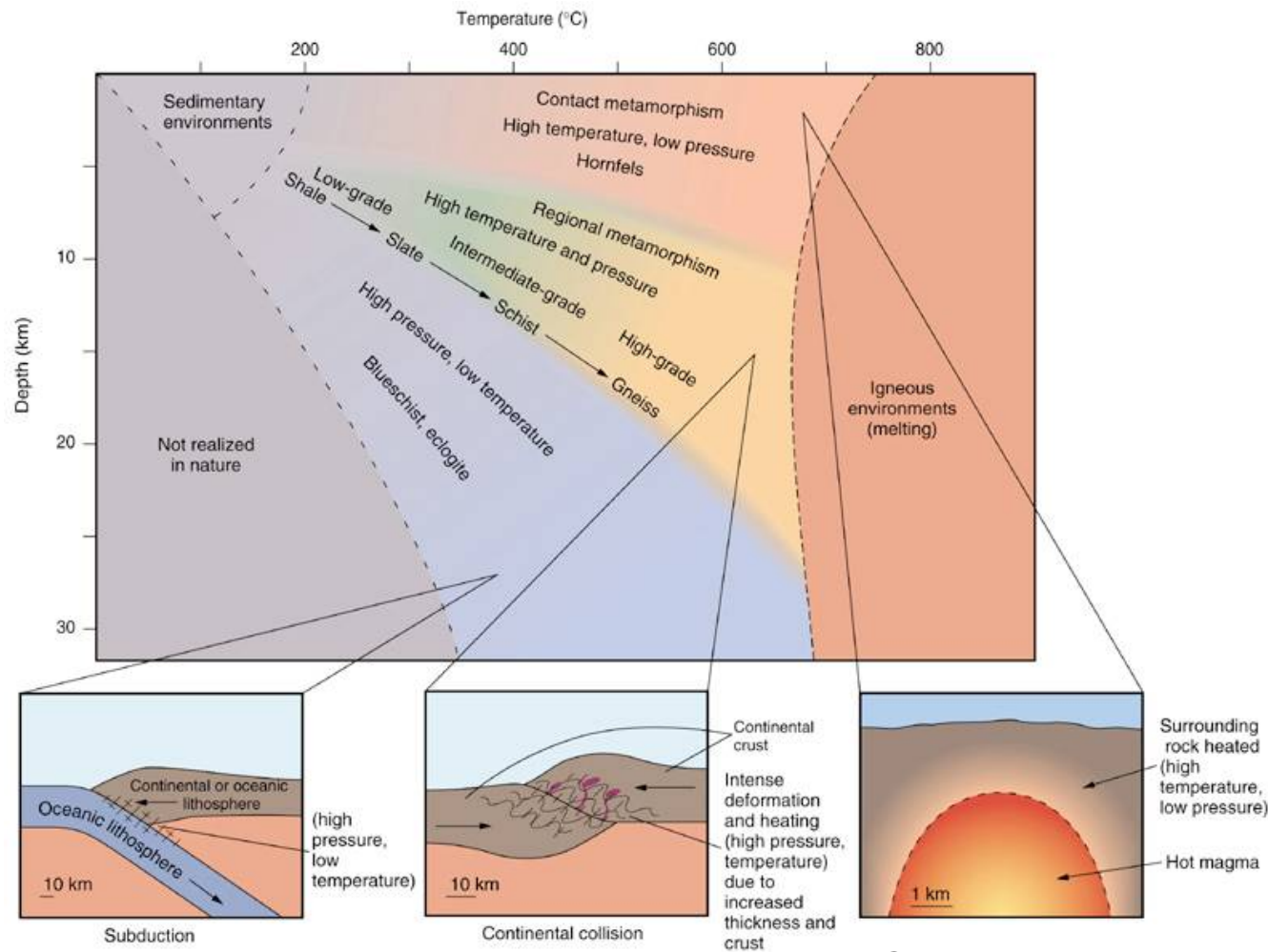
- Water acts as catalyst during metamorphism
 - Conduit for ion exchange in the growth of crystals
- Sources of fluid:
 - From pore fluids of sedimentary rocks
 - From cooling of plutons
 - Dehydration by heat of minerals (clay, micas)
- Aid in recrystallization of existed minerals

Types Metamorphism

Scale	Type of metamorphism	Metamorphic agents	Change in chemical composition
Regional	Regional Sedimentary rocks only: Burial	Heat and P	No
Local	Contact	Heat	
	Dynamic	P	
	Impact	P	
	Hydrothermal	Fluids	Yes

- Regional: large scale, plate tectonics
- Local: smaller scale, faults, igneous bodies

Types Metamorphism



Regional

(High P, Low T)

Regional

(High P, High T)

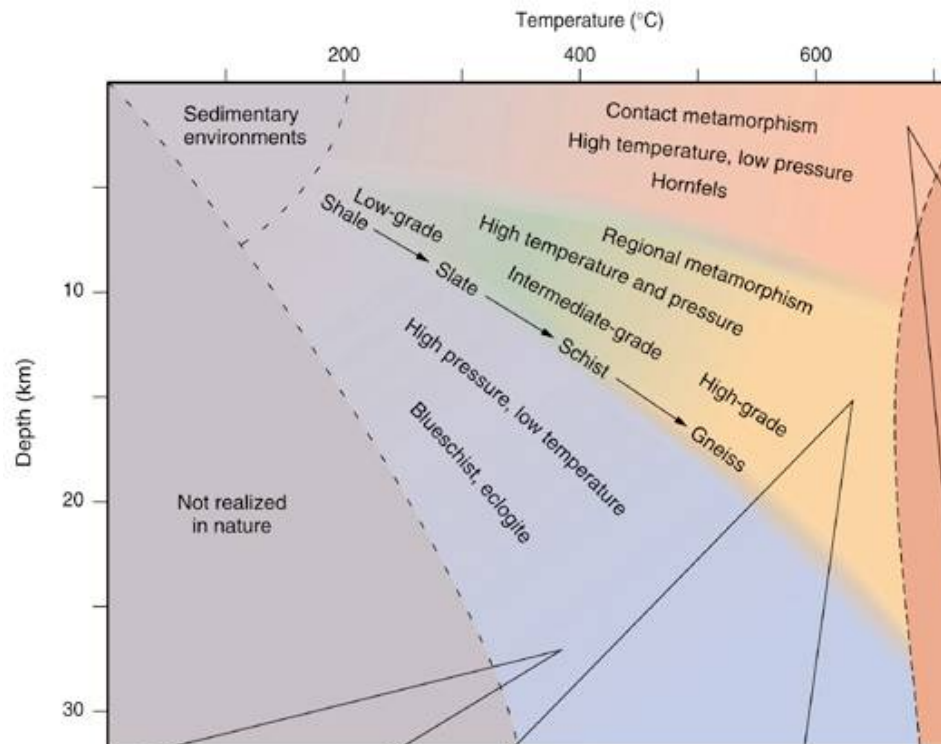
Contact

(Low P, High T)

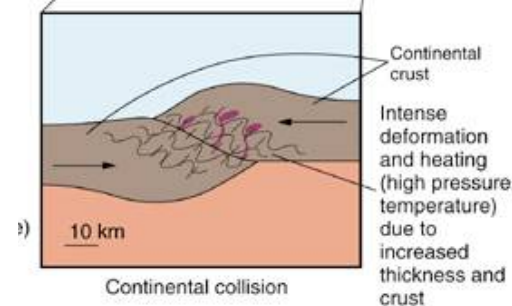
Regional Metamorphism

- Continental crust exposed to extreme T and P at convergent plate boundaries
 - Enormous thickness of rocks involved (9-12 km)
 - Greatest quantity of metamorphic rocks produced

Regional Metamorphism



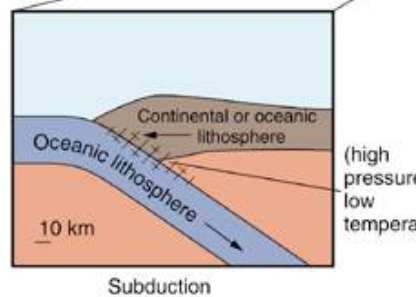
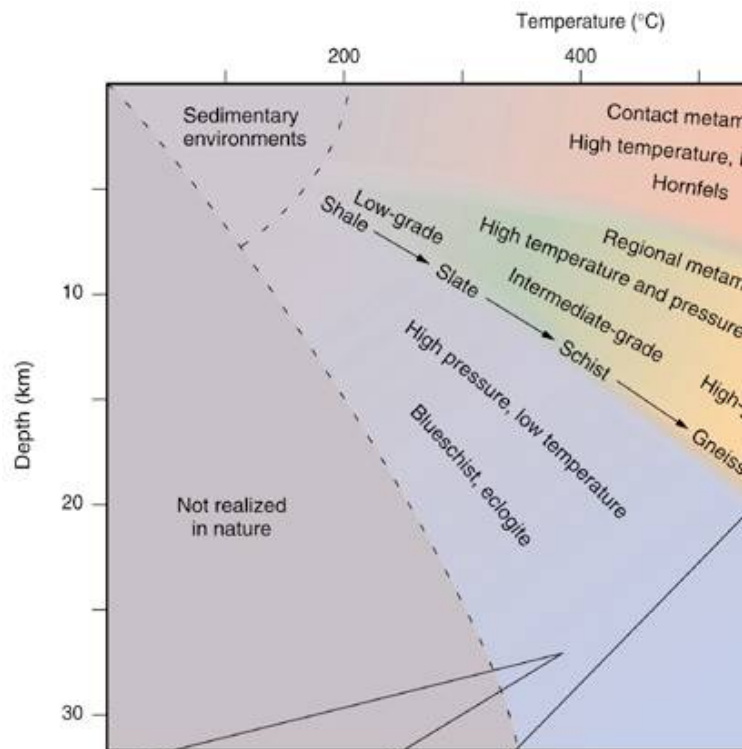
- Crustal thickening, continental collision
- Leads to foliation, alignment of minerals



Regional Metamorphism

- Low Grade: rock harder, more compact (interlocking crystals), minor increase in grain size
- Medium Grade: grain size increases, new mica crystals form
- High grade: grains recrystallize with preferred orientation (foliation)

Regional Metamorphism

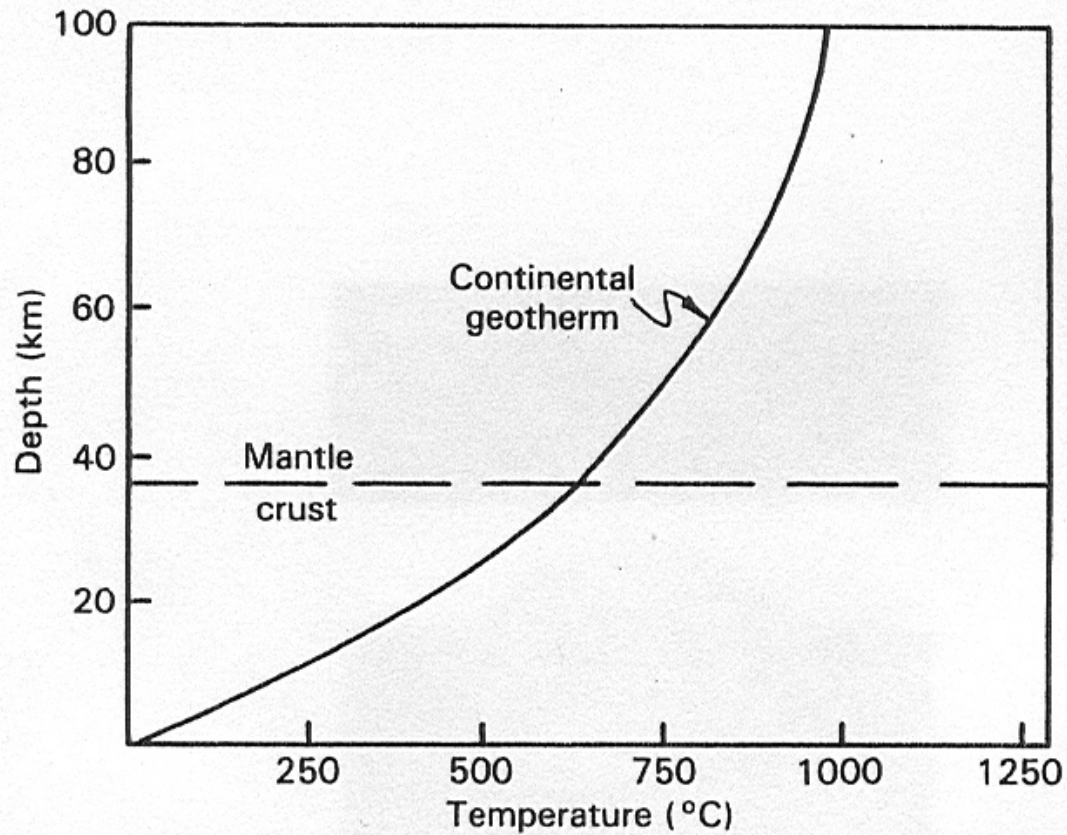


- At subduction zones of the lithospheric plate
 - Low T, High P
- Cold, slow to heat by conduction, but plunges quickly to great depth
- Sediments, basalts form high-P minerals

Burial Metamorphism

- Subset of regional metamorphism
 - Associated with thick sedimentary strata
- Metamorphism due to burial beneath other rocks
 - Certain mineral become unstable due to increased T and P, and change to new minerals
 - Required depth for metamorphism varies from one location to another depending on the **geothermal gradient**
- Bedding and other sedimentary structures preserved
- Clay becomes mica when buried several km's deep

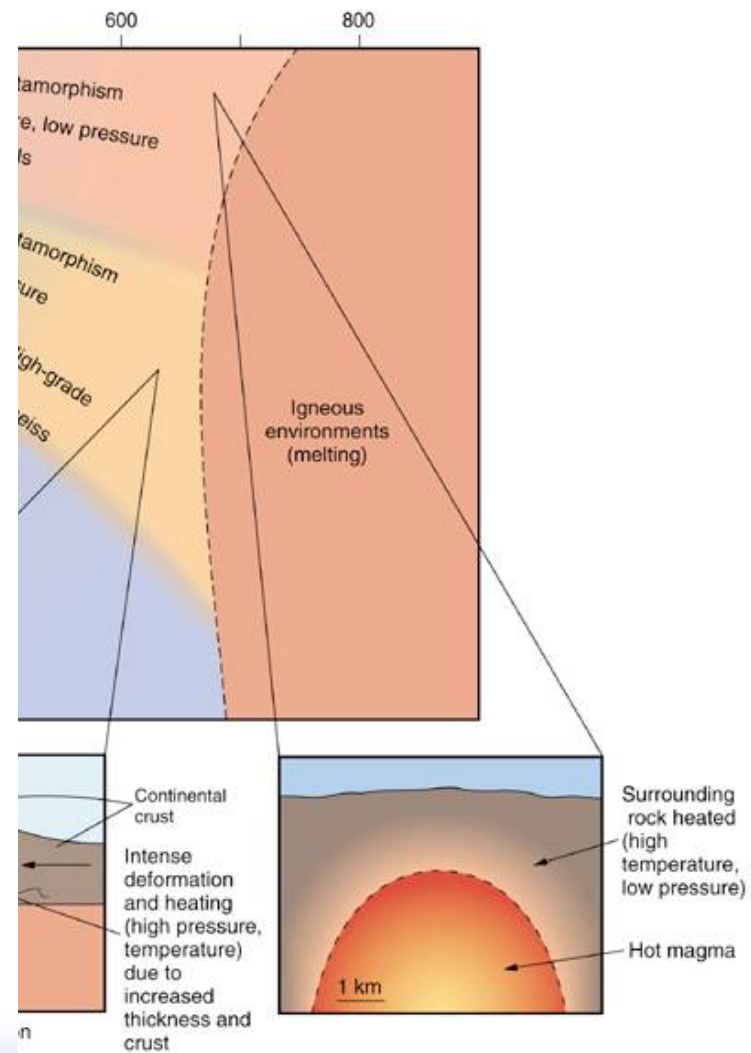
Continental geotherm



Ref.: Kehew, A.E. 1995. Geology for Engineers & Environmental Scientists. 2nd Edition. Fig. 5.2. Shown with permission.

Contact Metamorphism (Local)

- Magma intrudes a host rock leading to an increase in temperature
 - Increase in temperature leads to re-crystallization
- **Aureole**: zone of alteration, forms in the rock surrounding the igneous intrusion



Local Metamorphism: Contact

- **Aureole**
 - Extent depends on the size of the intrusion and the host rock
 - Sills, dikes → cm–m wide aureole
 - Batholiths → km wide aureole
- Typically fine-grained, dense and tough
 - No directional alignment of minerals because pressure is not a major factor

Dynamic Metamorphism (Local)

- Dynamic metamorphism is associated with localized events along fault zones
 - Rocks are broken and pulverized as rock blocks along fault zone grind past each other
 - **Fault breccia**: loose rock consisting of broken and crushed rock fragments
- Pressure is the only metamorphic agent

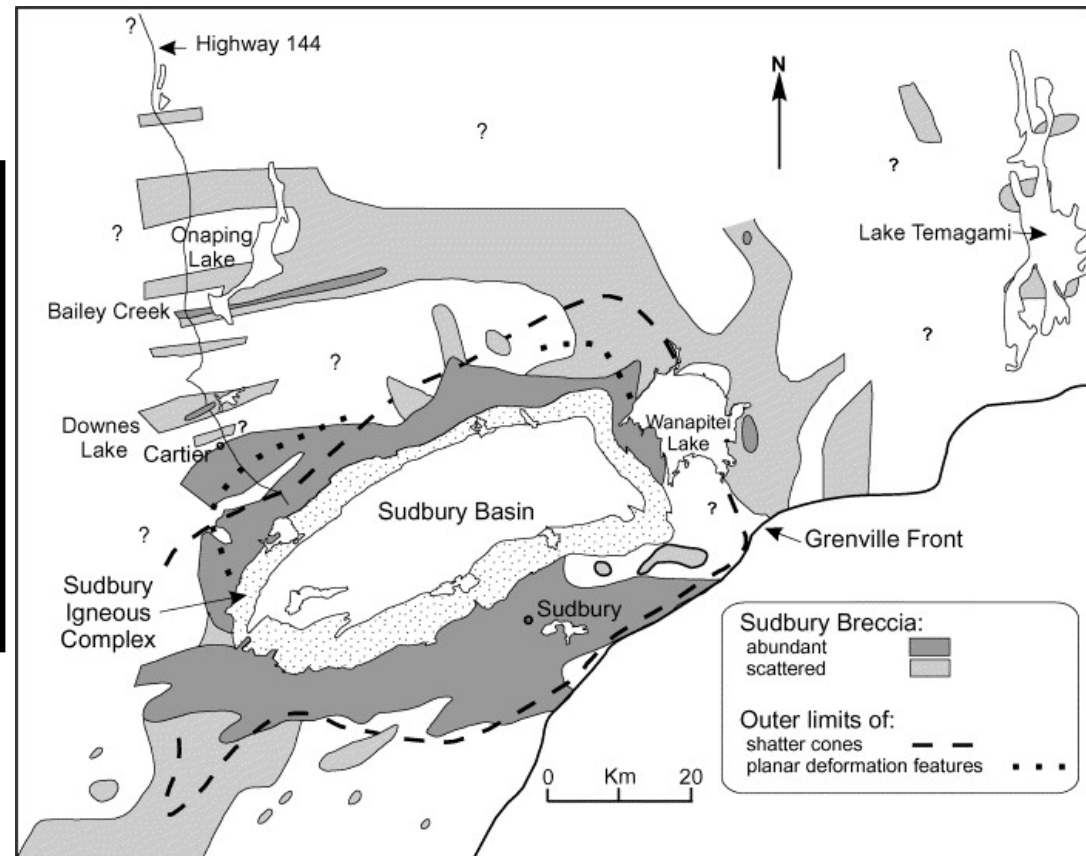
Impact Metamorphism (Local)

- Occurs when large space bodies strike the Earth's surface
 - Creates shattered and melted rocks adjacent to the impact crater
 - Shatter cone: a conical fragment of host rock that is formed from the high pressure of meteorite impact and has striations radiating from the apex
 - 2-30 GPa of pressure required

Blast shatter cone



Photo: W. Carter, Carleton U.



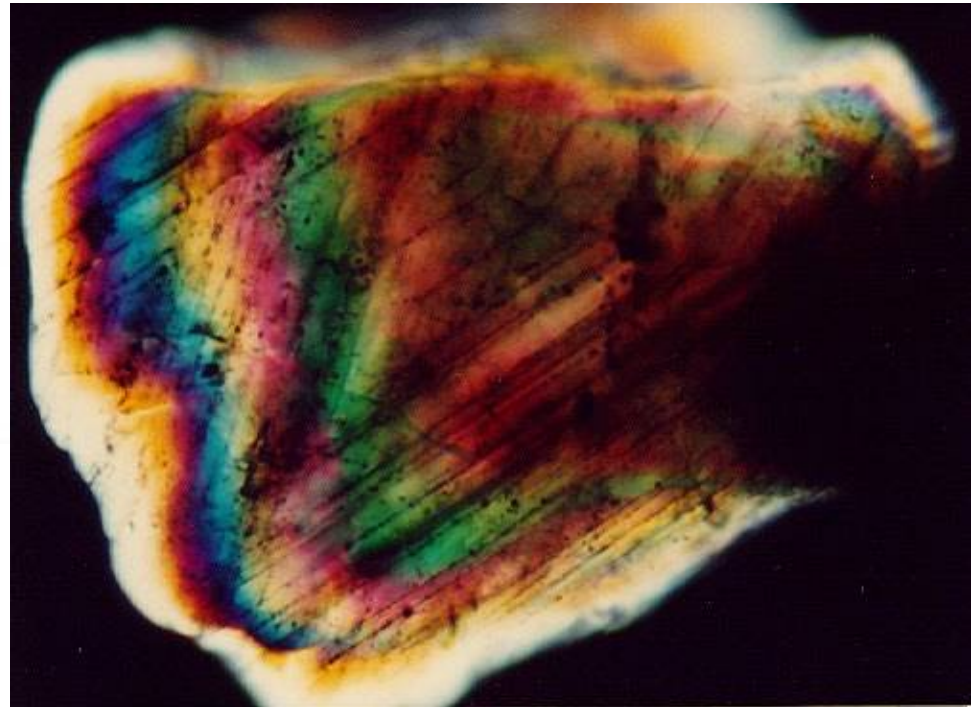
Rousell et al. 2003

[http://dx.doi.org/10.1016/S0012-8252\(02\)00091-0](http://dx.doi.org/10.1016/S0012-8252(02)00091-0)

Shocked quartz



Chicxulub crater, Mexico
dia > 180km



Dislocation of crystal orientations due to
shock wave (~5–8 Gpa)

Hydrothermal Metamorphism (Local)

- A form of **metasomatism**: change in the composition of a rock or mineral by the addition or replacement of chemicals
 - Chemical alteration caused by hot, ion-rich fluids (**hydrothermal solutions**) that circulate through cracks
 - Hydrothermal solutions are expelled from cooling plutons
 - Invade country rock

Local Metamorphism: Hydrothermal

- Add ions to rock to form new minerals
- May generate metallic ore deposits
 - Copper, gold, molybdenum, tin, iron
- Example:
 - Hydrothermal metamorphism is widespread along mid-ocean ridges
 - Fe-Mg rich minerals metamorphose into talc, serpentine

Classification

- Key terms, textures, characteristics, naming convention

Protolith

- **Protolith**: original rock from which a given metamorphic rock was formed
 - Most metamorphic rocks have the same overall chemical composition as their protolith
 - Original mineral makeup determines, to large extent, the degree of which each metamorphic agent will cause change
- Example: quartz sandstone is the protolith of quartzite

Protolith

- Different metamorphic minerals can be produced from the same original rock under different temperature and pressure conditions
- Example: shale → slate → schist → gneiss

Metamorphic textures

- Three metamorphic textures:
 - Foliated textures
 - Pressure is the key metamorphic agent
 - Mostly associated with regional metamorphism
 - Non-foliated texture
 - Temperature is the key metamorphic agent
 - Mostly associated with contact metamorphism
 - Porphyroblastic texture
 - Large crystals growing in a finer matrix
 - Example: garnet

Foliation

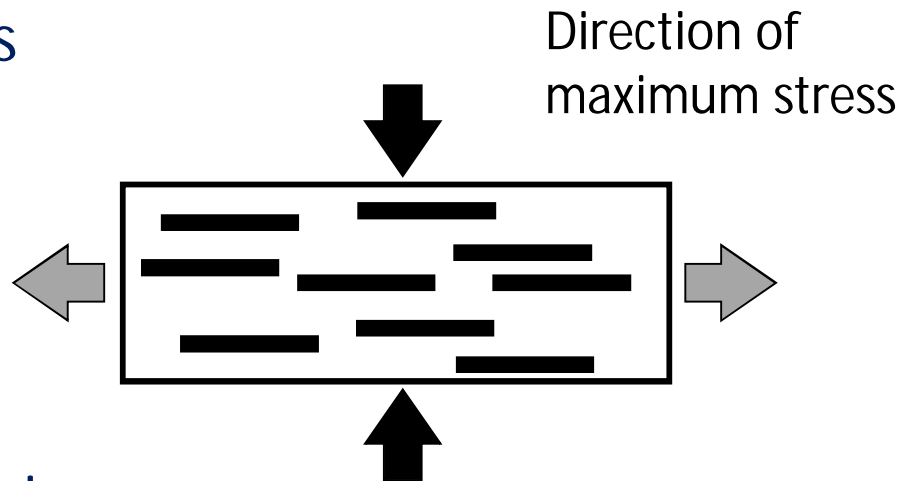
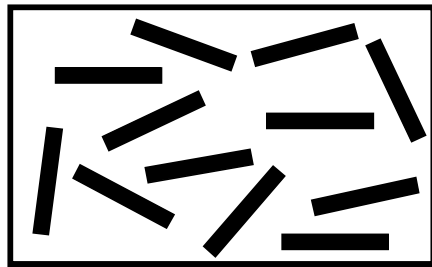
- **Foliation**: parallel orientation of mineral grains within a metamorphic rock
 - Foliation give metamorphic rocks a banded appearance
 - Watch out! Foliation resembles the bedding of sedimentary rocks
 - .. The two are produced by different mechanisms

Foliation

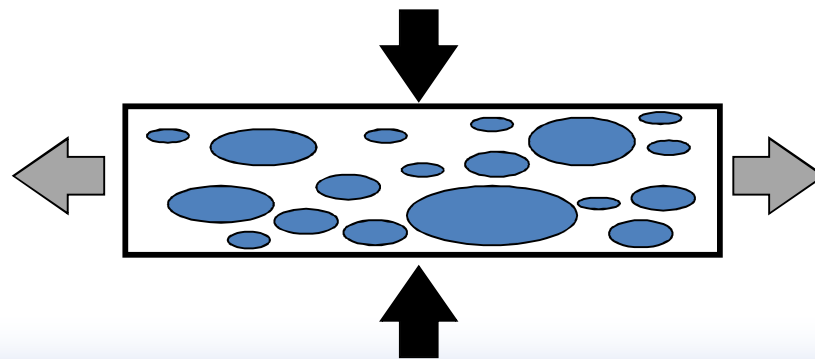
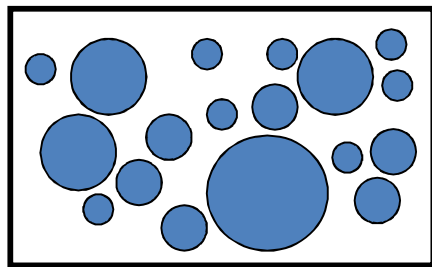
- Foliation can develop in different ways:
 1. Rotation of platy minerals under pressure
 2. Recrystallization of minerals under pressure
 3. Flattening of rounded grains in the direction of maximum stress
- New mineral orientation is perpendicular to the direction of maximum stress

Foliation

- Rotation of platy minerals



- Flattening of rounded grains



Foliation

- Different expressions of foliation
 - Parallel alignment
 - Compositional banding
 - Zones of weakness where rocks can easily split into thin sheets

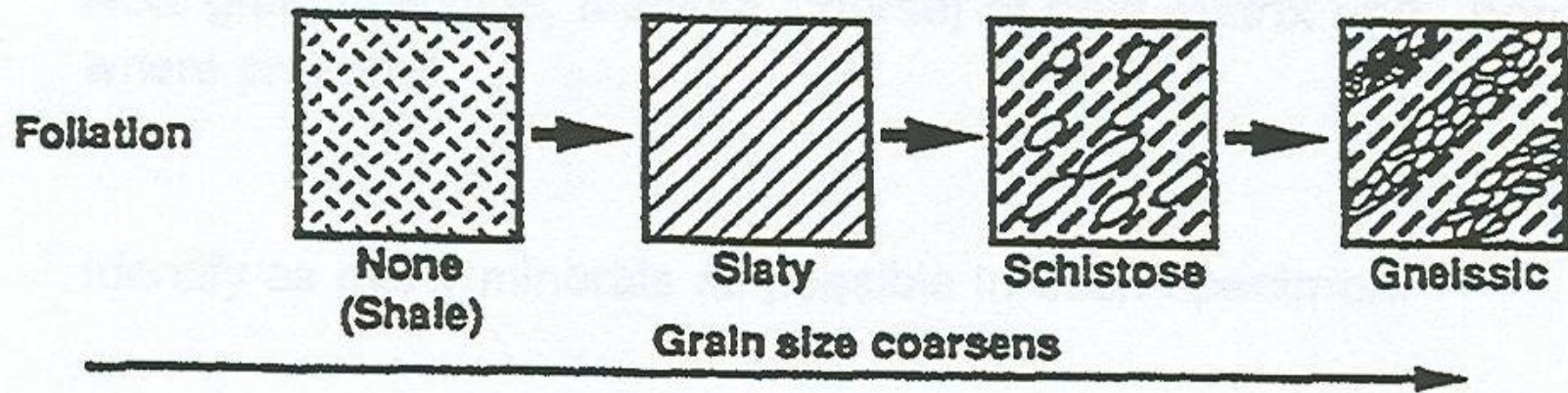
Foliated textures

- **Slaty cleavage**
 - Closely spaced planar surfaces along which rocks split
- **Schistose**
 - Platy mineral grains grow larger when rock is subjected to higher T and pressure
 - Rock exhibit a “scaly” appearance
 - Platy minerals visible to unaided eye

Foliated textures

- **Gneissic texture**
 - During higher grades of metamorphism, ion migration results in segregation of minerals
 - Distinct banded appearance due to segregation of dark and light silicate minerals

Foliated textures



Increasing Metamorphic grade

Non-Foliated textures

- Develop in environments where deformation is minimal
 - Mostly associated with contact metamorphism
- Develop from rocks composed primarily of one mineral, exhibiting equi-dimensional crystals
- Examples:
 - Limestone → marble
 - Quartz sandstone → quartzite

Porphyroblasts

- Large mineral grains/crystals (porphyroblasts) surrounded by a fine-grained matrix of other minerals
- Grows within the solid state
- Examples: garnet, staurolite, andalusite

Porphyroblasts

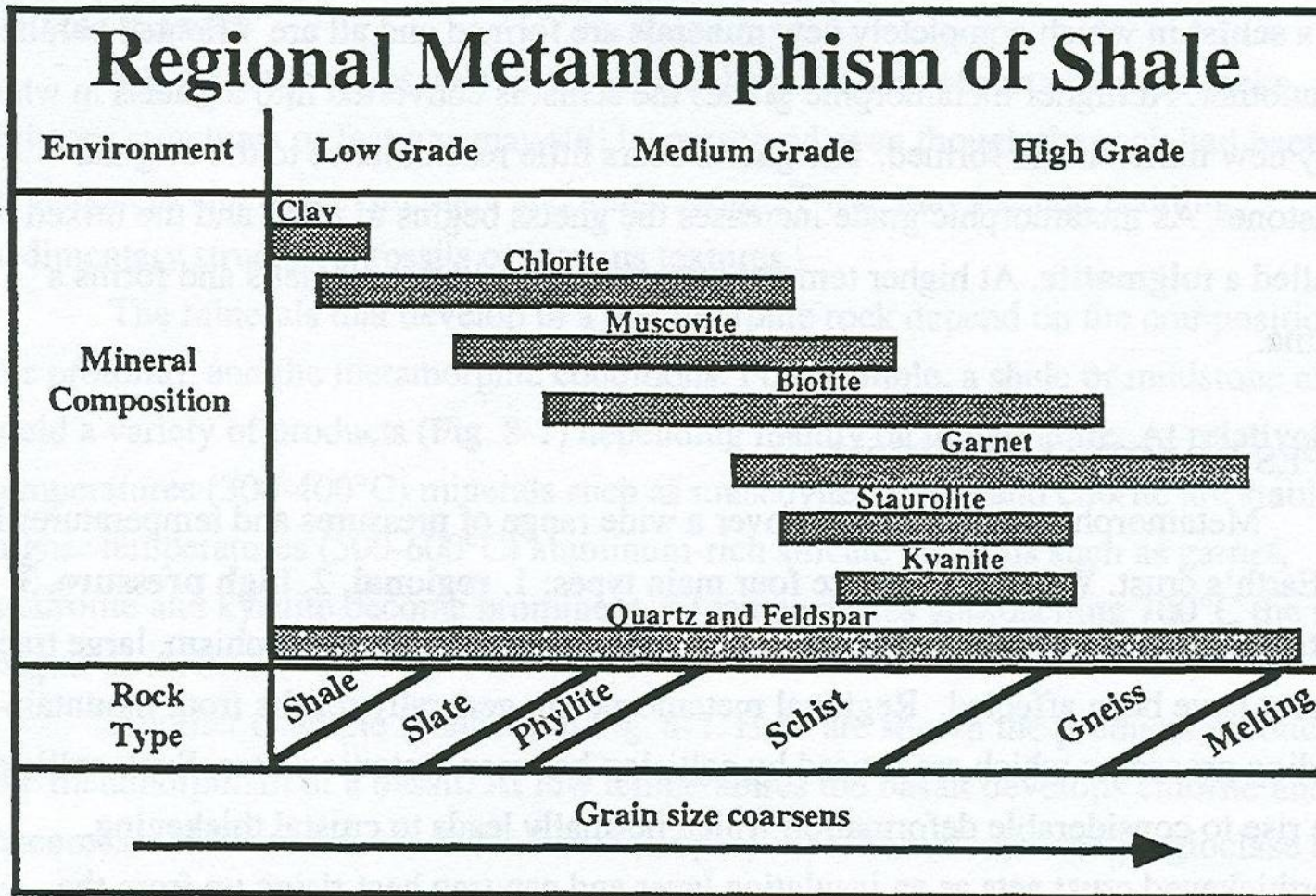


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Classification

- Metamorphic rocks are classified according to:
 - Texture
 - Protolith

Metamorphism of Shale

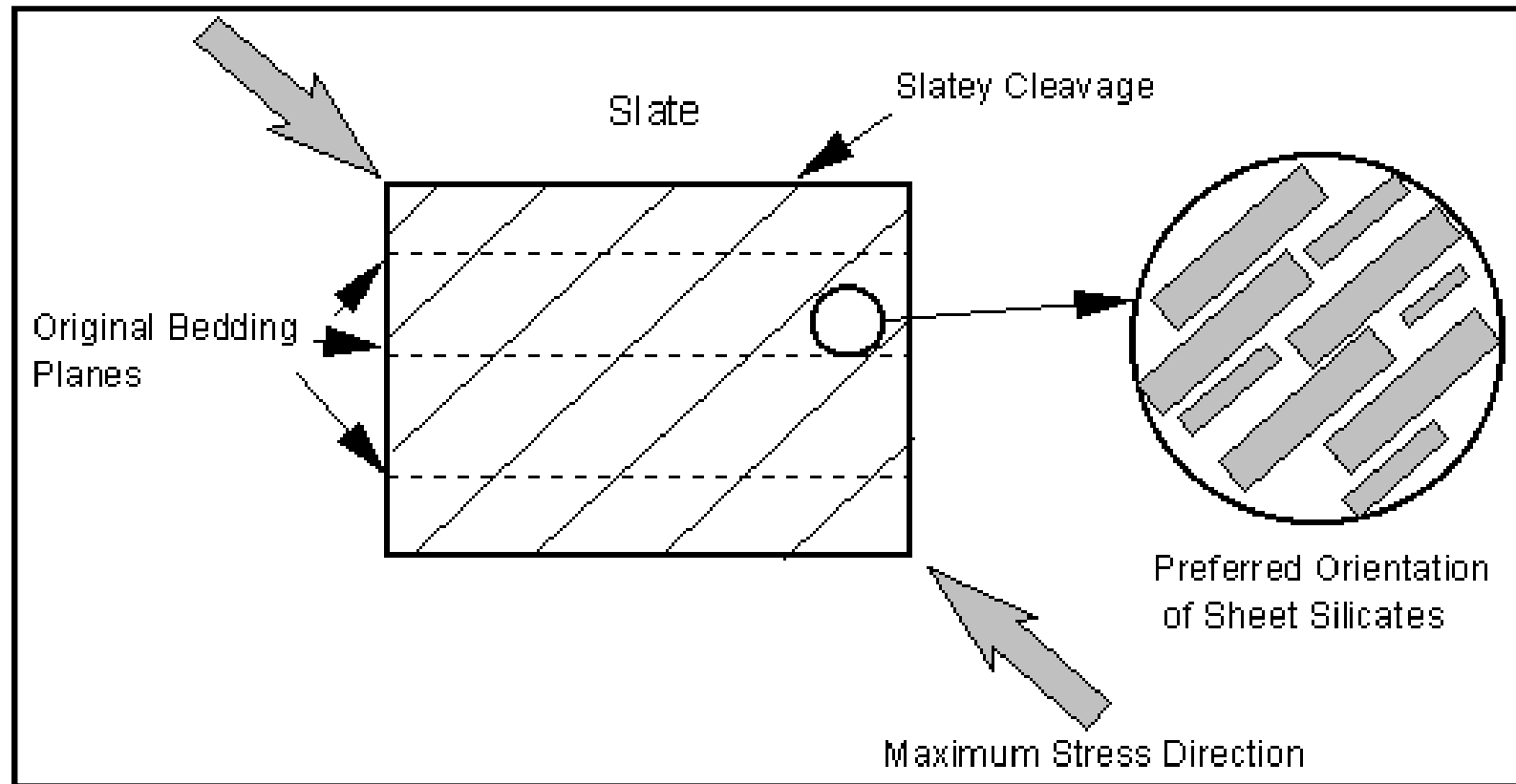


Ref.: EARTH2404 Lab manual

Slate

- Generated from low-grade metamorphism of shale
 - Very fine grained (minute mica flakes)
 - Color depends on mineral constituent
 - Black: organic material
 - Red: iron oxide
 - Green: chlorite minerals
- Excellent cleavage
 - Tiles
- Original bedding may or not be preserved
 - Slaty cleavage may develop at an angle to original bedding

Slate



Source: Earth Science Australia
(<http://earthsci.org>)

Slate



Shale



Slaty cleavage

Phyllite

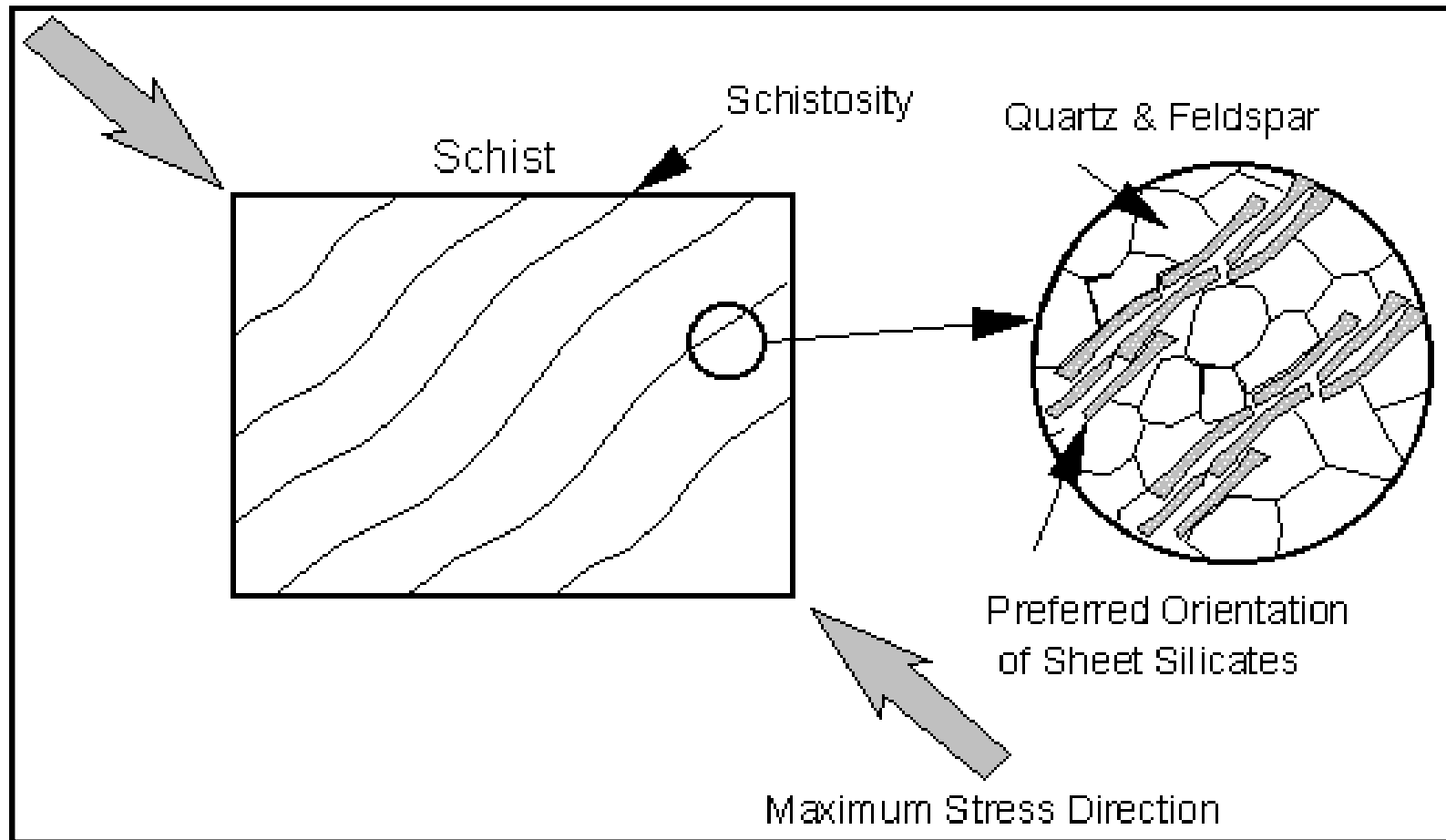
- Lies between slate and schist
 - Characteristic rock "sheen"



Schist

- Generated from intermediate-grade metamorphism of shale
 - Often related to mountain building
- Medium to coarse grained rock readily split into thin scales
 - Platy mica minerals dominate
 - Often contains accessory minerals unique to metamorphic rocks (e.g. garnet)
- Term “schist” describes the texture
 - To indicate composition, mineral names are used (e.g. mica schist)

Schist



Source: Earth Science Australia
(<http://earthsci.org>)

Schist

- Naming convention applies porphyroblast(s) first
 - If there are more than one, name most abundant first



Garnet schist

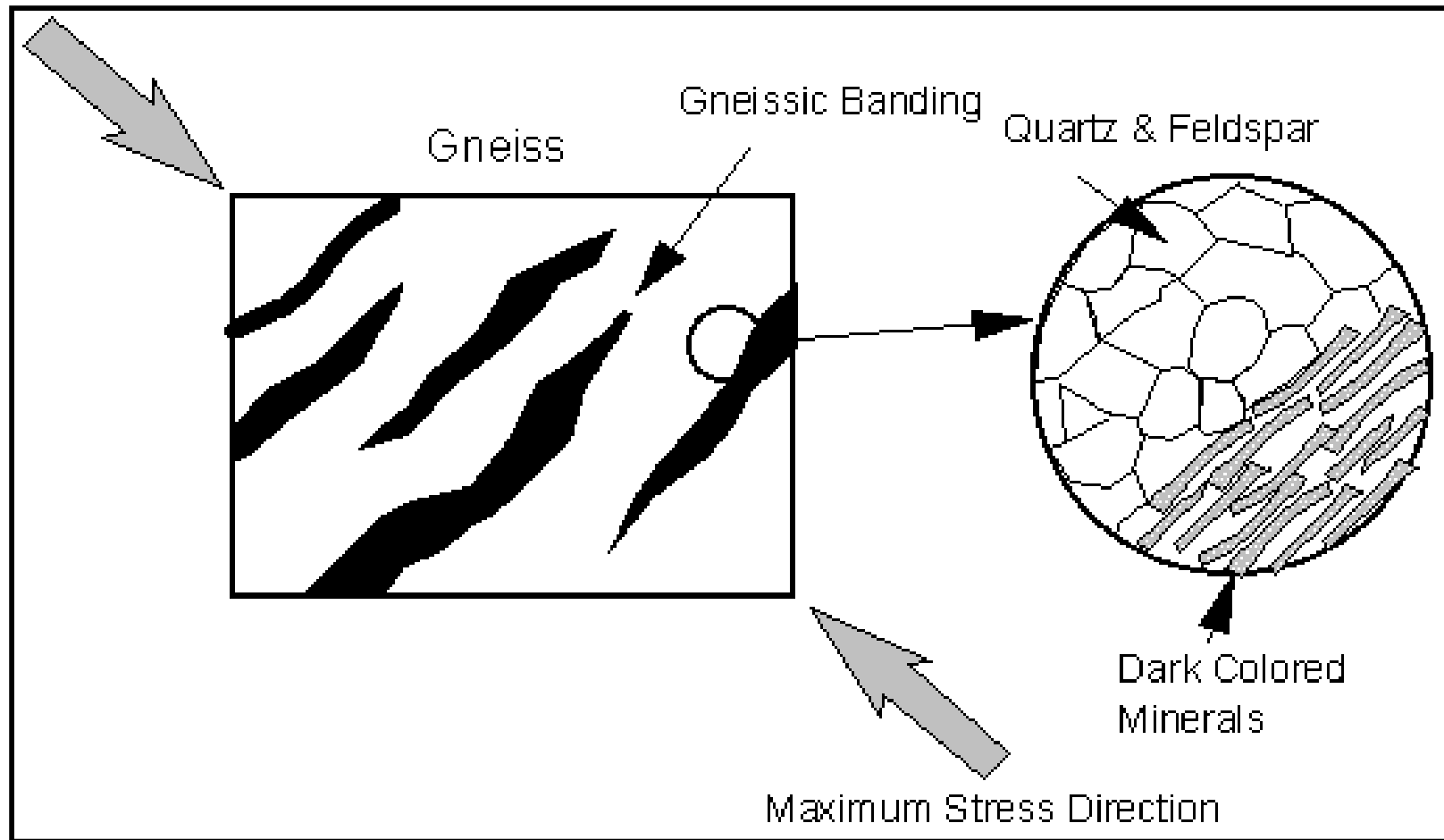


Muscovite schist

Gneiss

- Generated from high grade metamorphism of shale or granite
- Medium to coarse grained rock with elongated granular minerals (quartz, feldspar, hornblende) and a lesser amount of platy minerals (mica)
- Banded appearance
- Break in irregular fashion

Gneiss



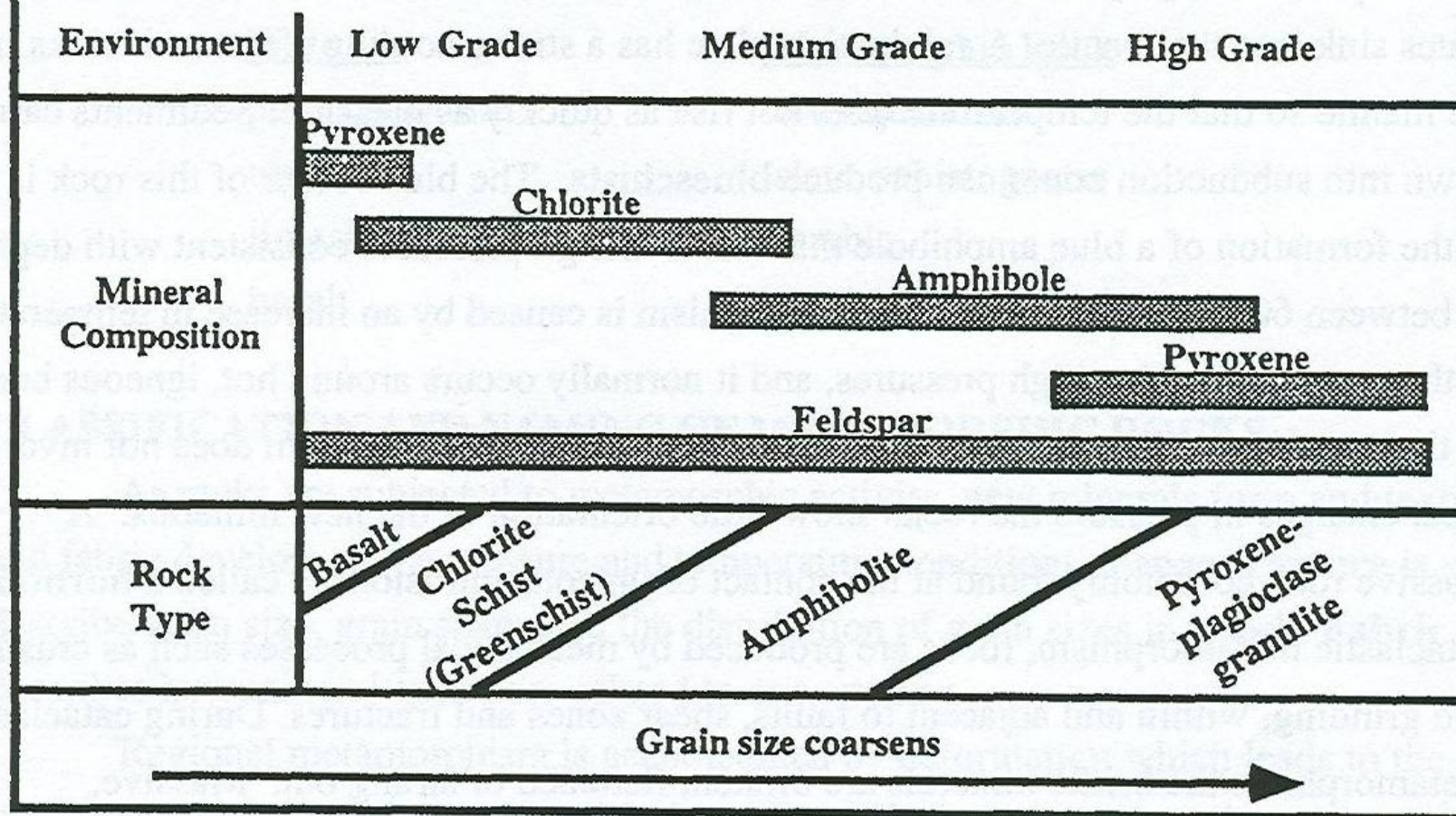
Source: Earth Science Australia
(<http://earthsci.org>)

Gneiss



← Acasta Gneiss of the Slave Province of the Shield (4.03 Ga old)

Regional Metamorphism of Basalt



Marble

- Common non-foliated metamorphic rock
- Protolith is limestone or dolostone
- Composed essentially of calcite or dolomite
- Decorative stone



Quartzite

- Common non-foliated metamorphic rock
- Protolith is quartz sandstone
 - Quartz grains are fused together
- Very hard rock
 - Rock splits through quartz grains, not in between original grains



Engineering considerations

- Non-foliated metamorphic rocks have similar properties as intrusive igneous rocks
 - Great strength

Engineering considerations

- Foliated metamorphic rocks are susceptible to failure along specific planes
 - Orientation of foliation planes is a critical factor in slope stability
 - Avoid transferring load (from bridges, dams, building foundations) onto foliated rock mass in a direction parallel to the foliation

Engineering considerations

- Landslides associated with schist
 - Generally controlled by orientation of schistosity
 - Presence of slick minerals (e.g. talc, clay, graphite) may increase sliding potential

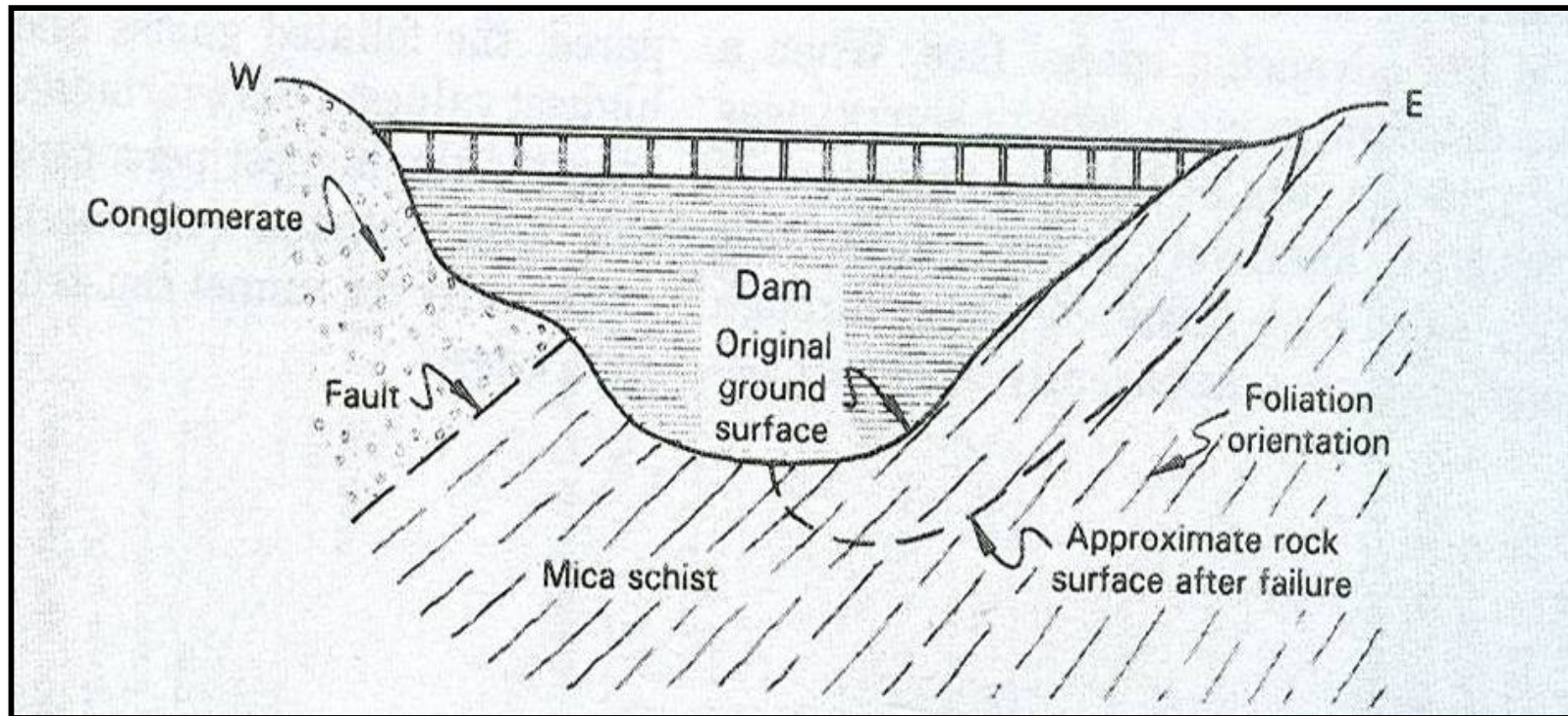
Failure of St. Francis Dam, CA

- https://www.youtube.com/watch?v=_6AvEZO34xl
- Built in 1926 to regulate the Los Angeles Aqueduct
- 12 March 1928, the dam suffered from catastrophic failure
 - Releasing 45 Billion liters
 - > 600 estimated deaths from the resultant flood

Failure of St. Francis Dam, CA

- Geological context
 - Poorly consolidated conglomerate of gravels, sands and muds held together by clay and cemented by gypsum
 - Schist with steeply dipping foliation planes
 - Load of the dam parallel to foliation

Failure of St. Francis Dam, CA



Ref.: Kehew, A.E. 1995. Geology for Engineers & Environmental Scientists. 2nd Edition. Fig. 5.16. Shown with permission.

Failure of St. Francis Dam, CA

- Water seeping from reservoir weakened rocks beneath the dam
 - Fault zone acts as a conduit for water
 - Degradation of weak bounds between mica grains along foliation planes in the schist
 - Water dissolved gypsum cement in the conglomerate

Next: Plate Tectonics!!

